

Water during planet formation and evolution

University of Zurich, 12–16 February 2018

Confirmed Speakers

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Magma dynamics and devolatilization of planetesimals during planet formation



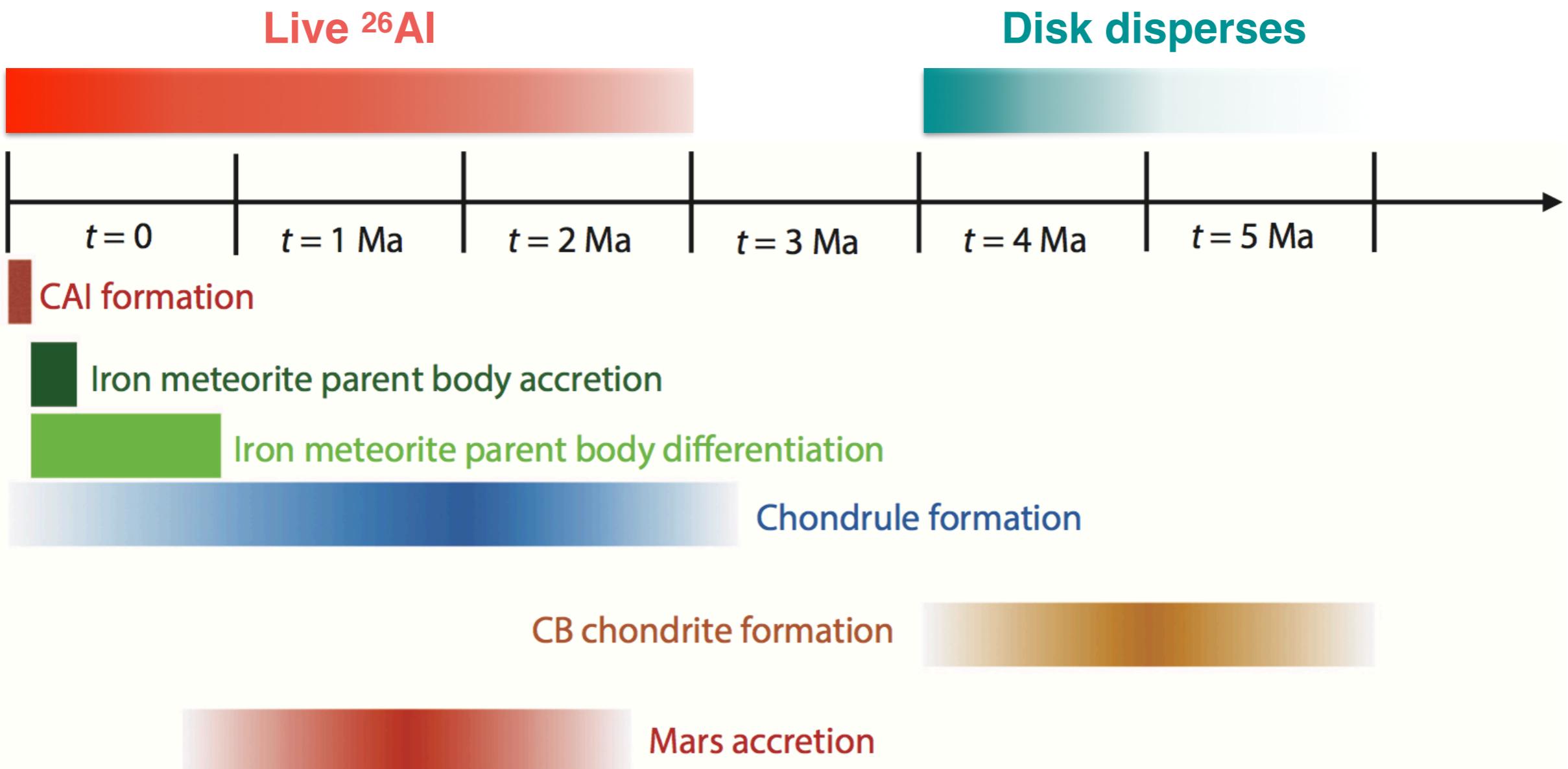
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Taras V. Gerya (ETH Zürich), Tobias Keller (Stanford),
Richard F. Katz (Oxford), Yann Alibert (CSH Bern)

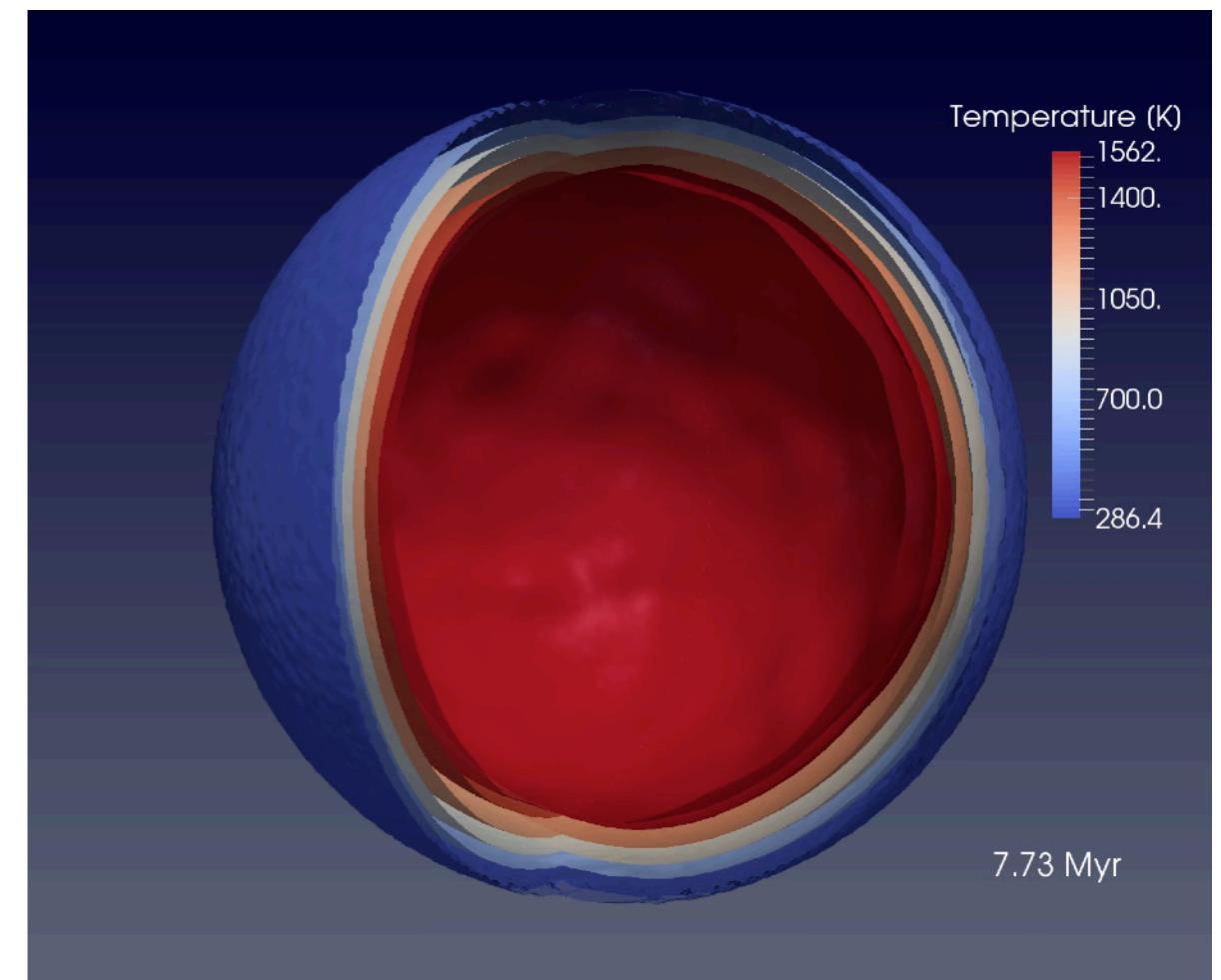
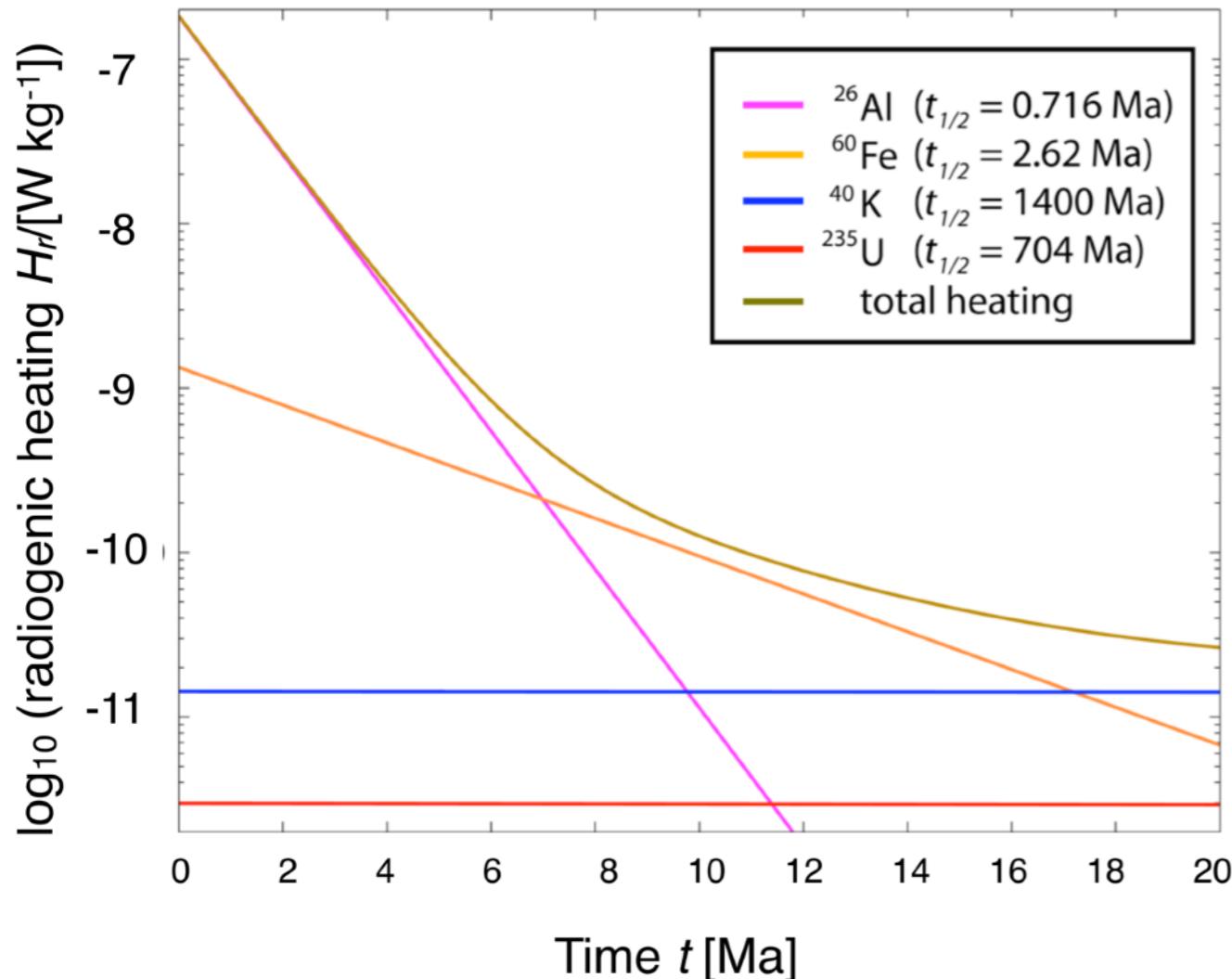


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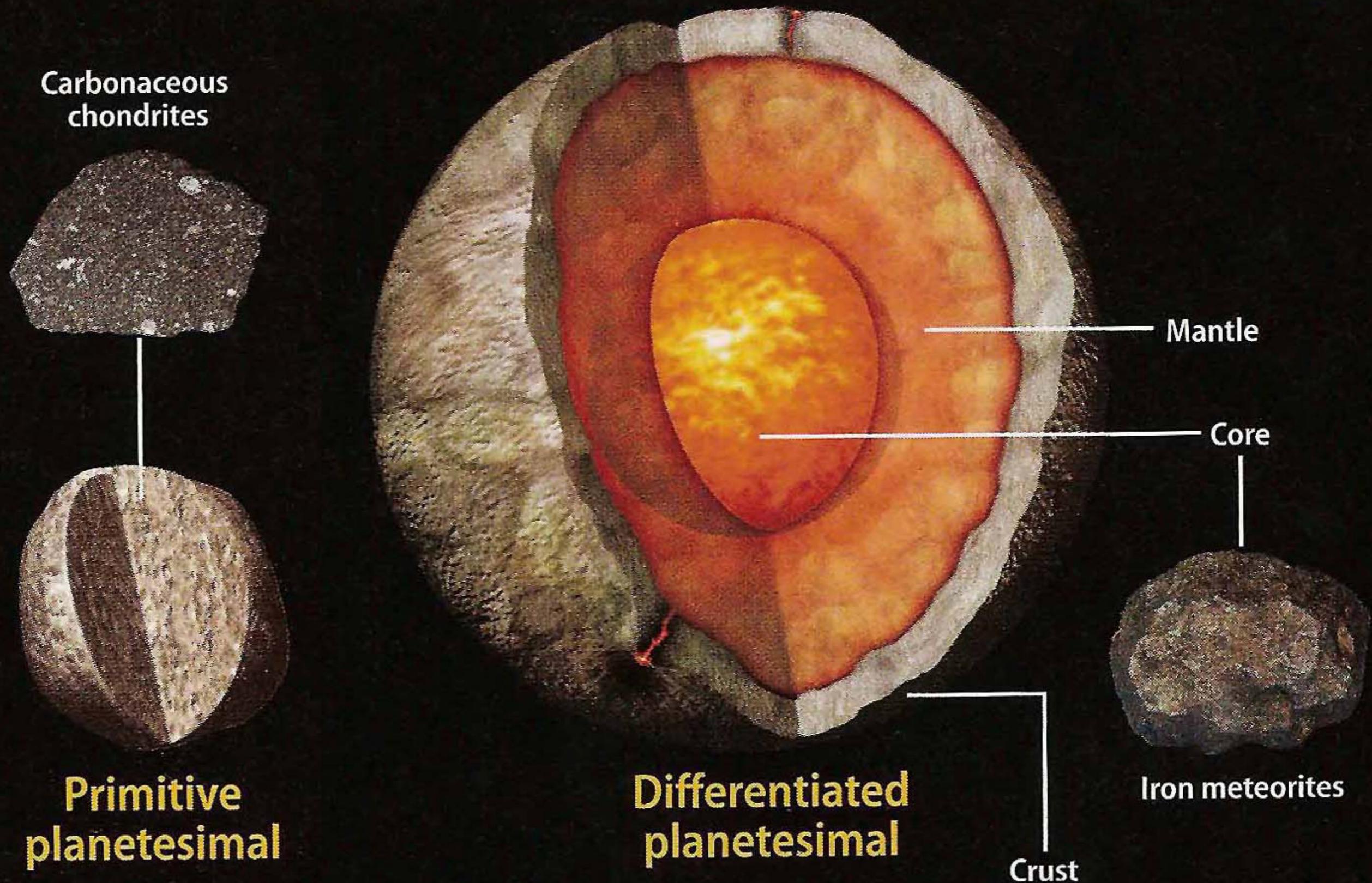
Planet formation: solar system materials



Planetesimal evolution regimes



Meteorite origins



Primitive
planetesimal

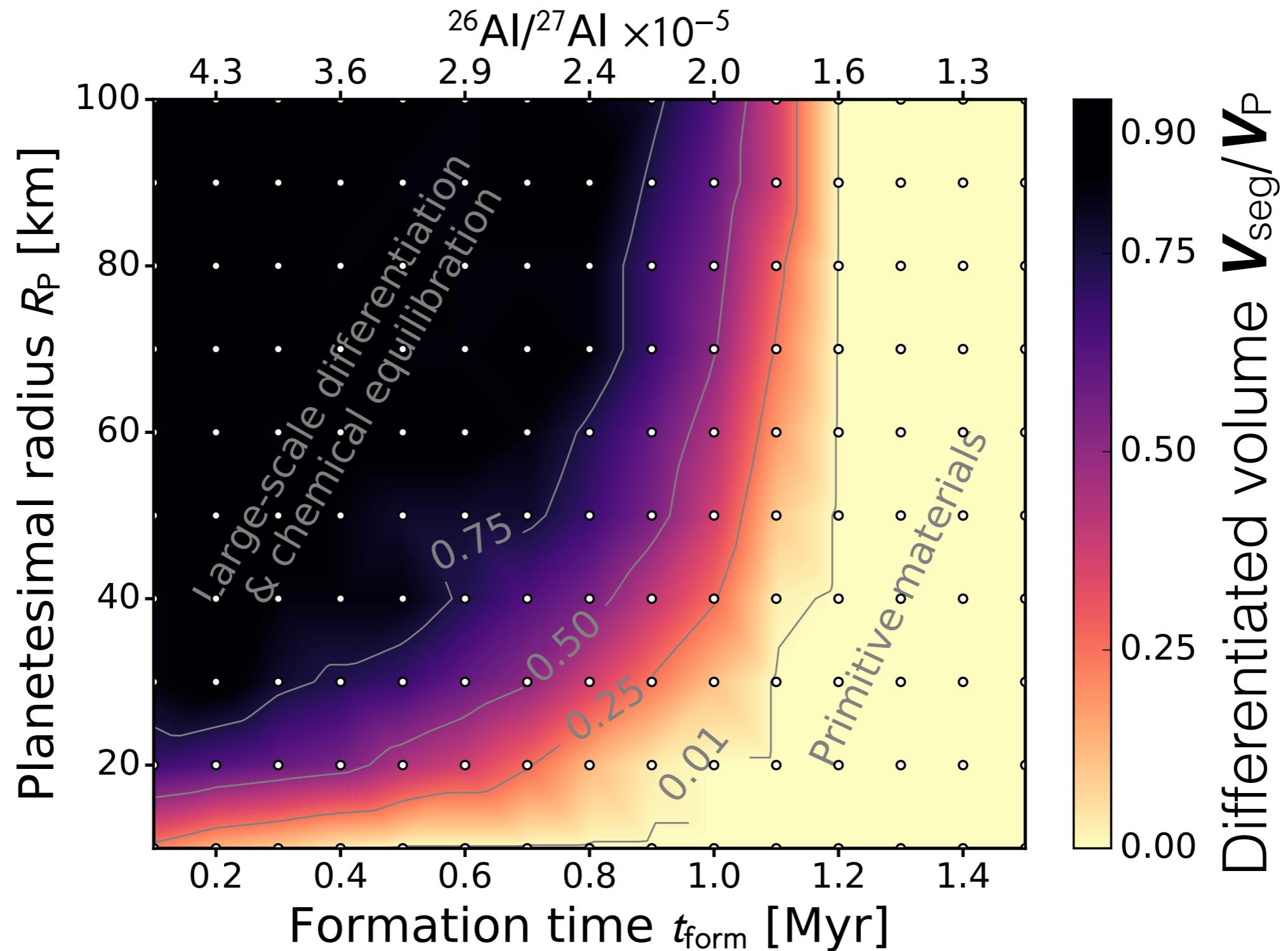
Differentiated
planetesimal

Crust

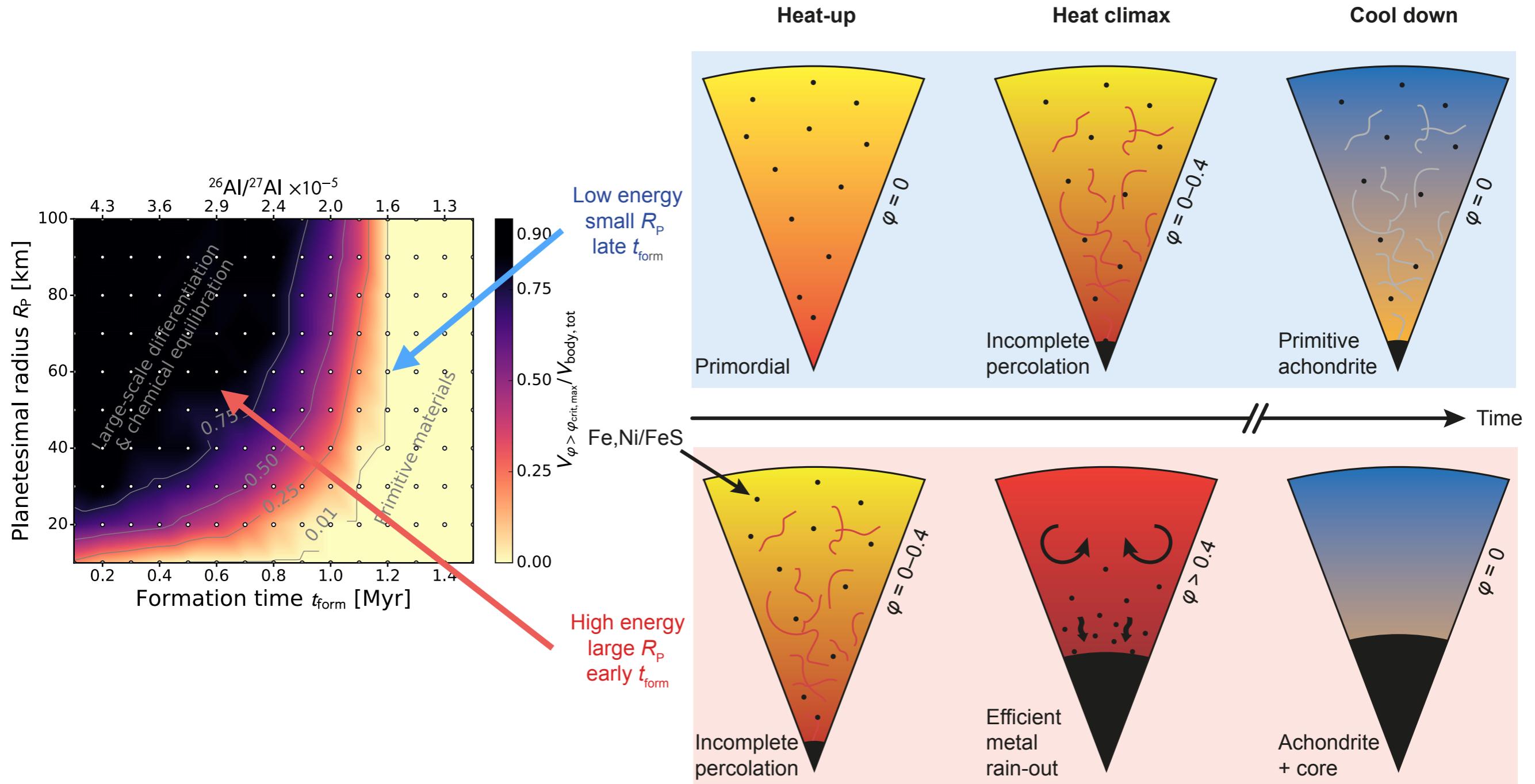
Iron meteorites

Carbonaceous
chondrites

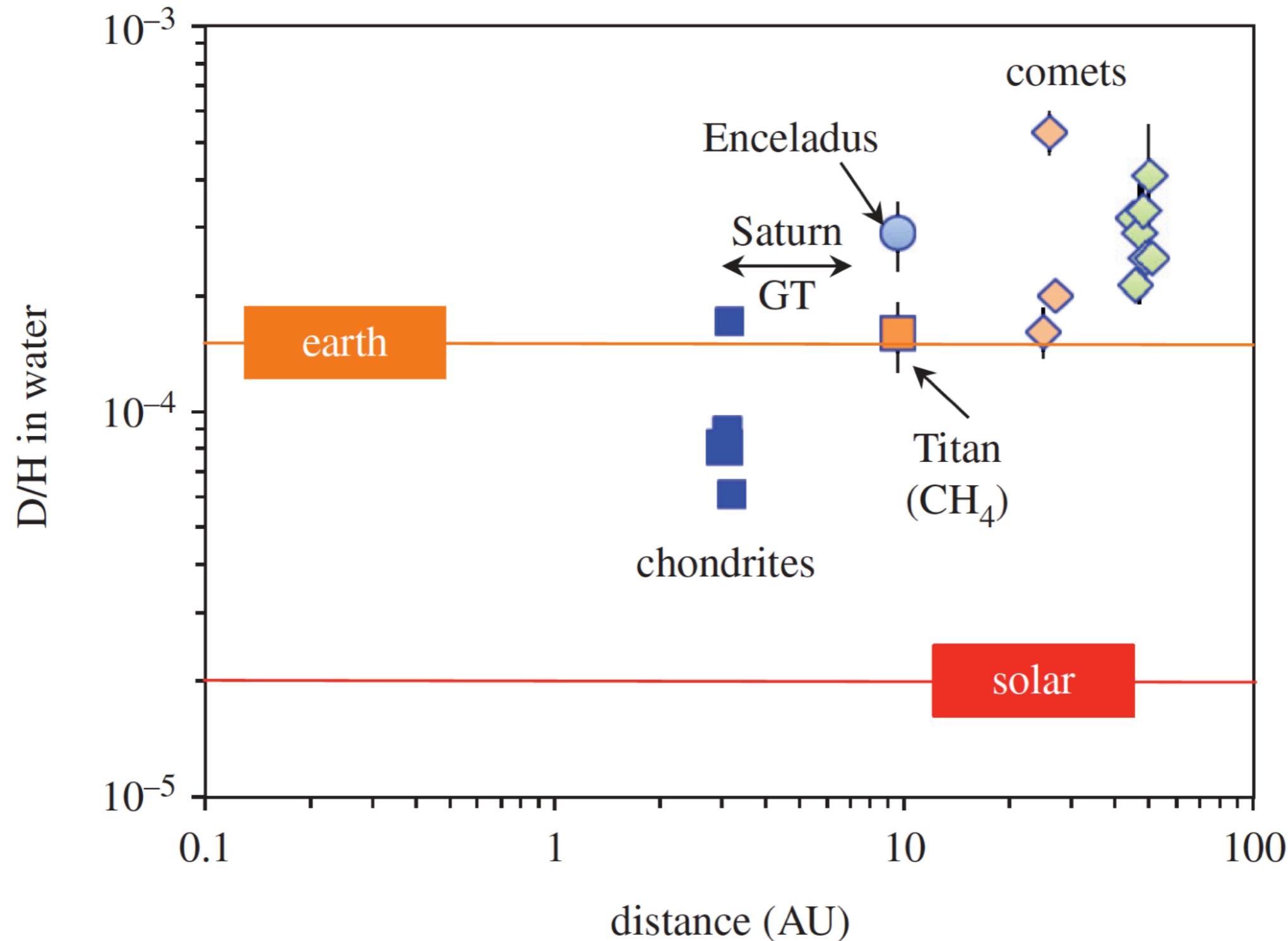
Thermo-mechanical interior evolution



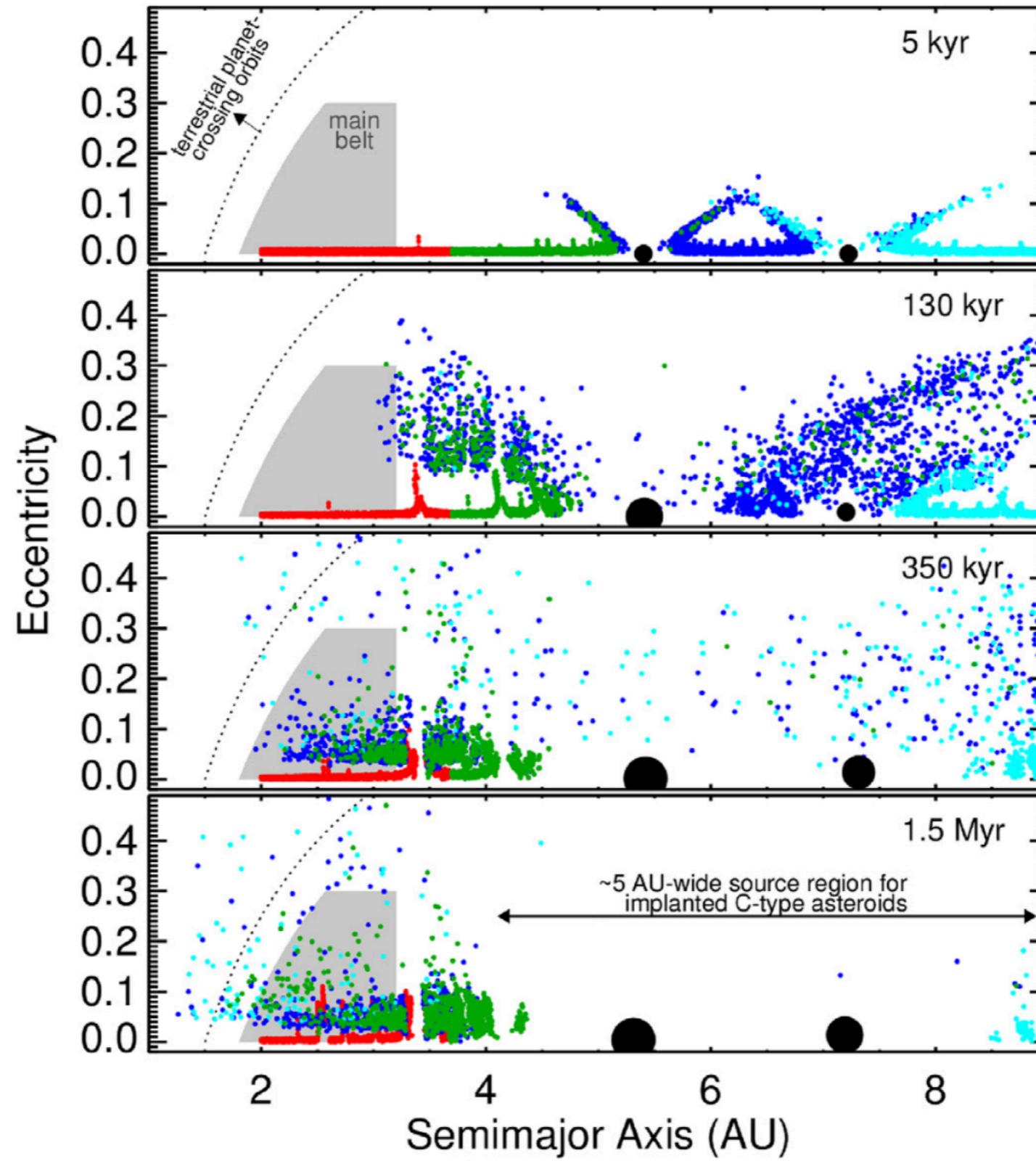
Thermo-mechanical interior evolution



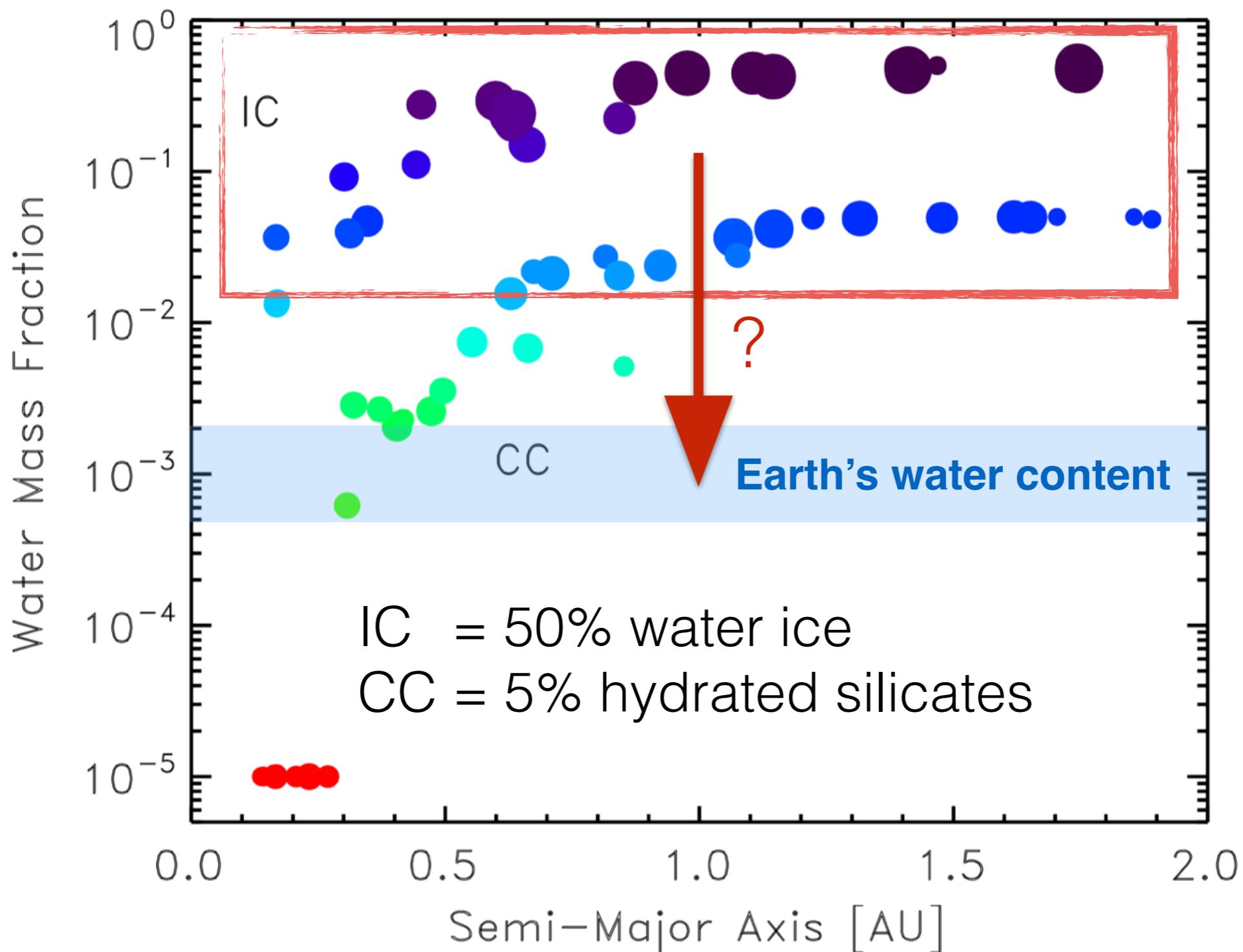
Delivery of Earth's water



Water delivery during accretion



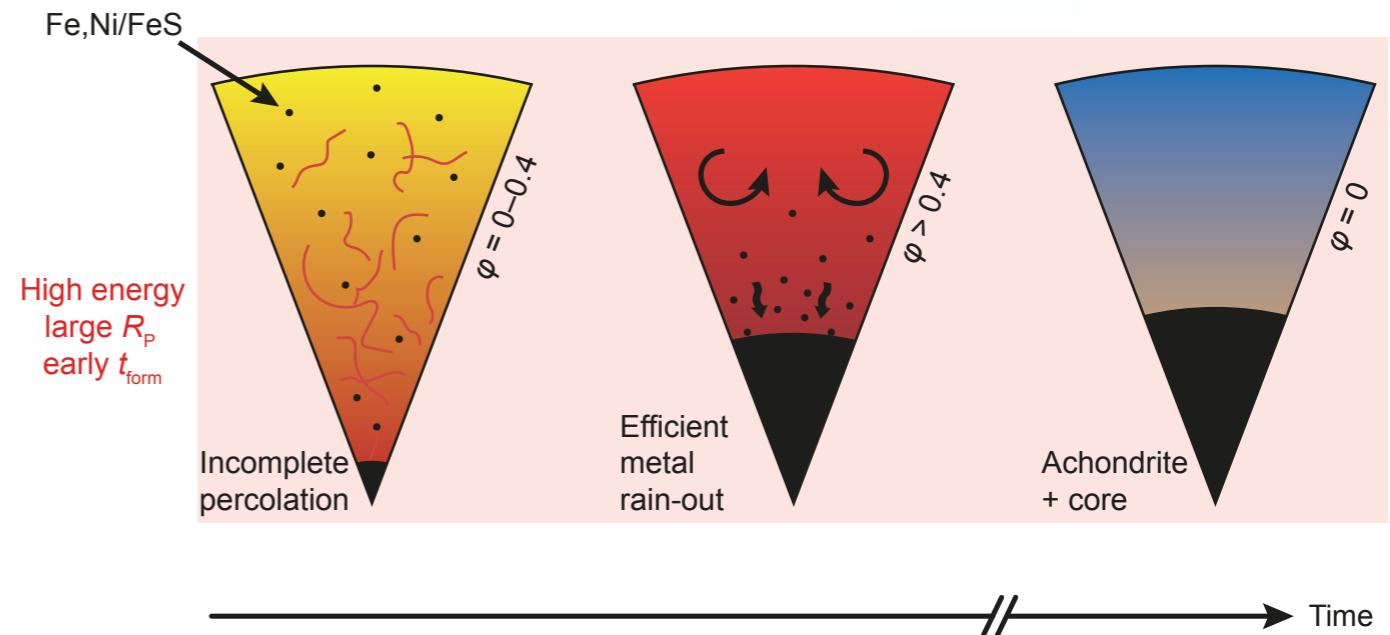
Fine-tuning Earth-like water abundances



Consequences?

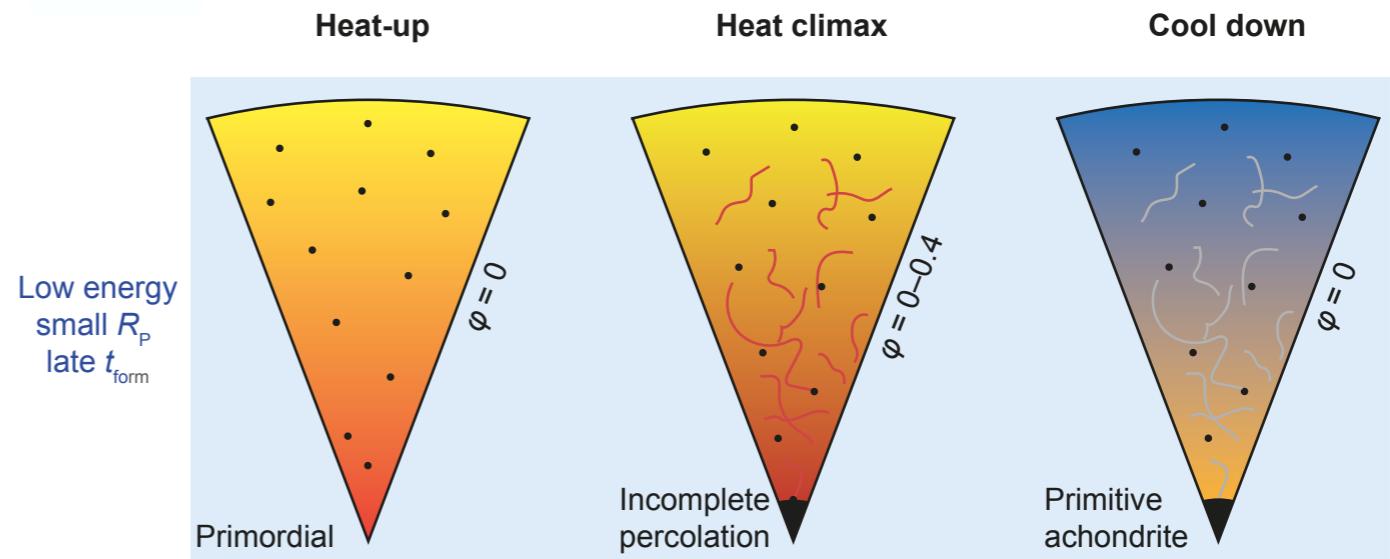
1. High energy: silicate melting & differentiation

- Radiogenic heating,
connection to meteorites
- Magma-rock overturn?

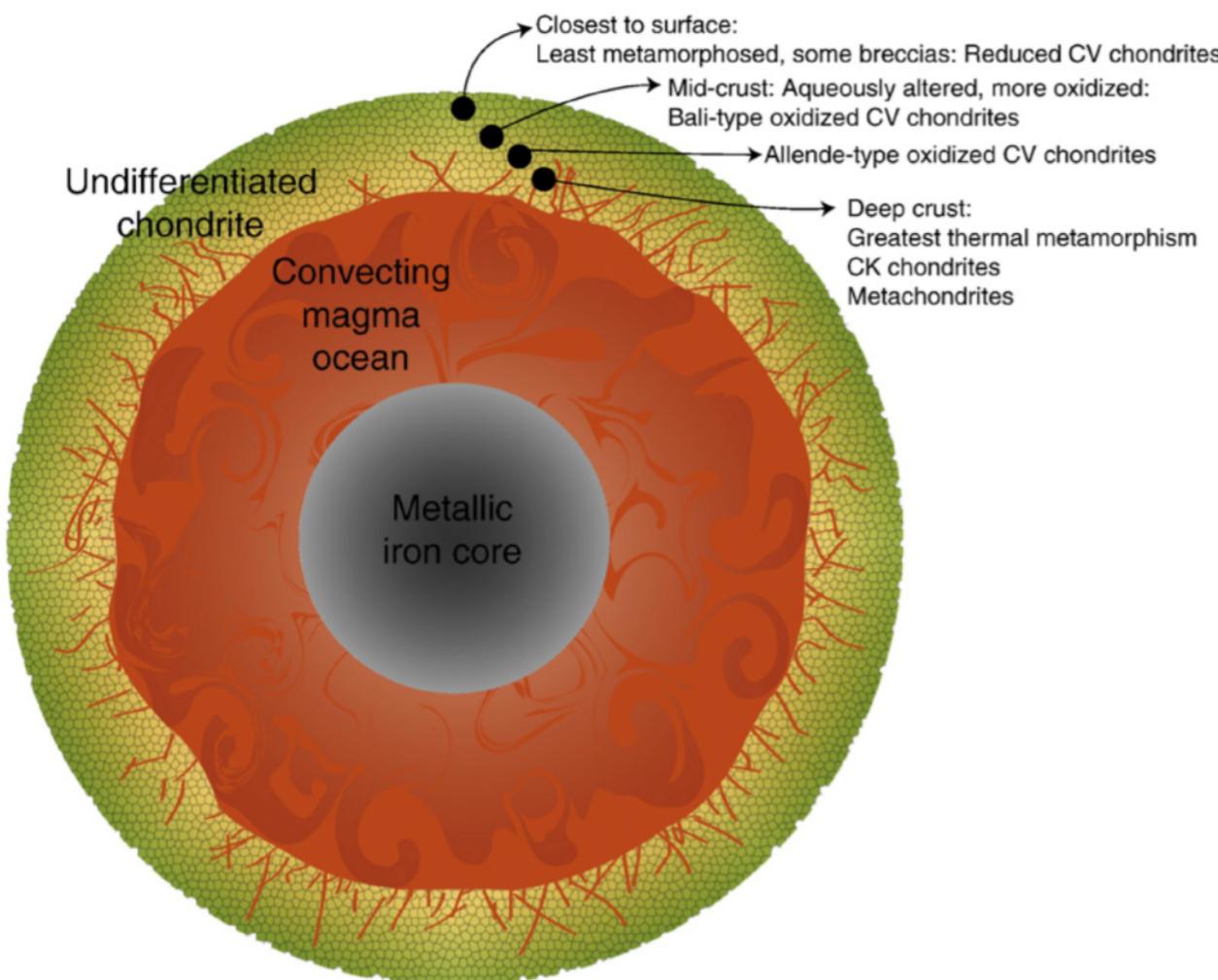


2. Low energy: fluid flow and volatile outgassing

- Dehydration of planetary materials?
- Water delivery to terrestrial planets?



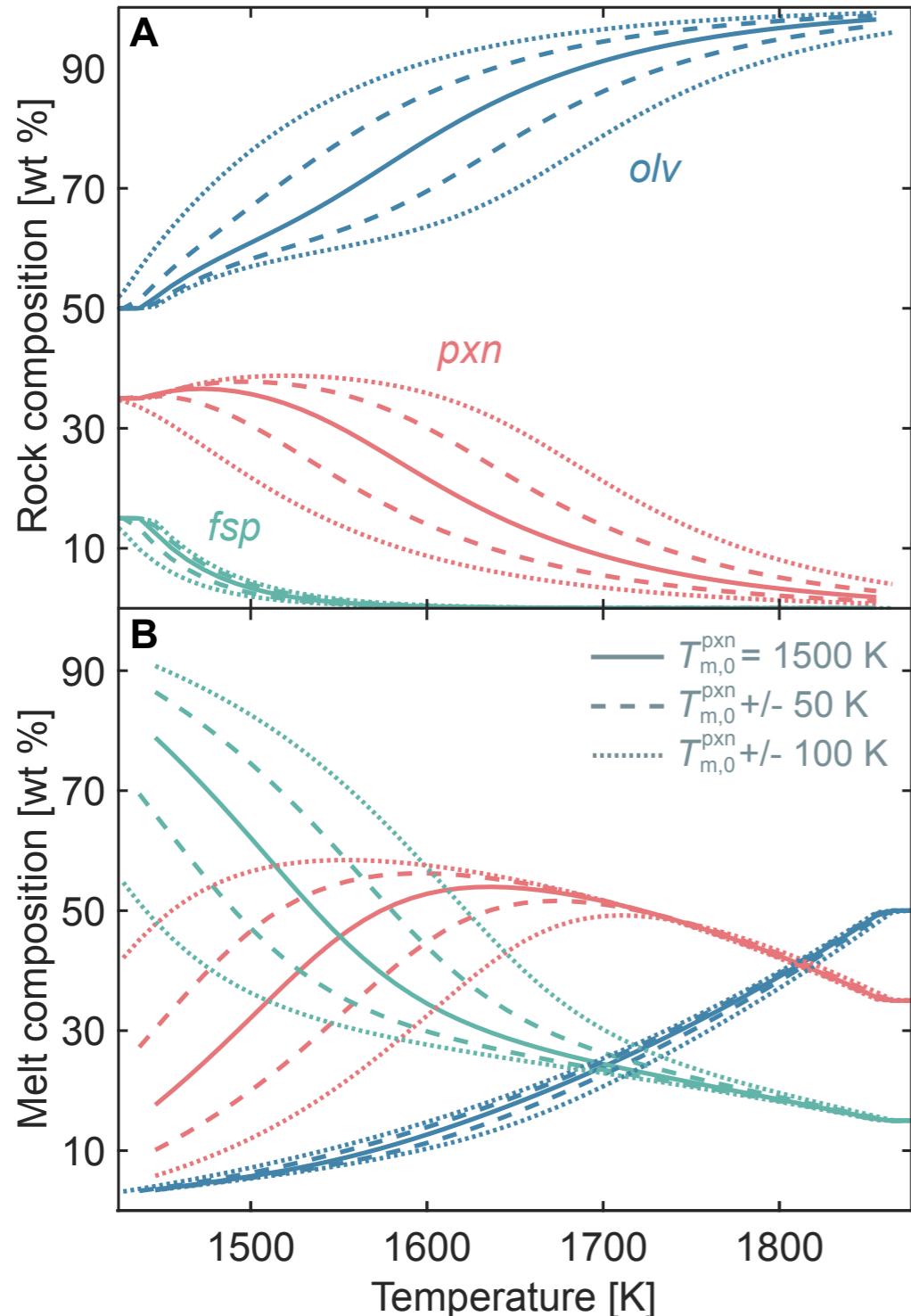
Part I: Not so ‘onion’ after all?



The ‘onion shell’ model

- Thermal evolution usually assume that melt and solid **do not separate**
- However: **Silicate melt** may be **buoyant** relative to planetesimal mantle
 - Planetesimals stripped of their heat source? Magma oceans at all?
 - Primordial crust preserved? Internally differentiated objects among C-complex-like asteroids?
—> (21) Lutetia

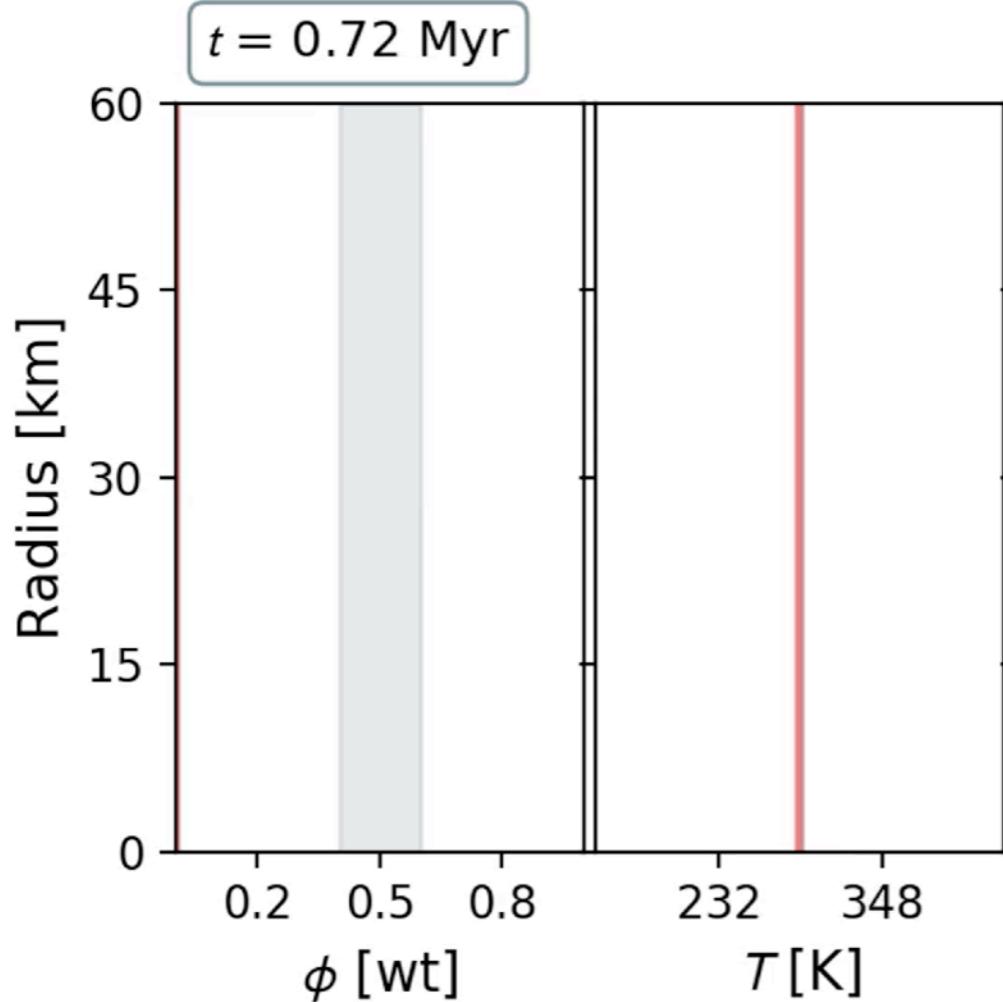
Thermo-chemical two-phase model



- Split up planetesimal rock body into multiple components, follow individually
- **Two-phase, thermo-chemical** evolution in 1D-column models (R_{DMC} method)
- ‘Dry’ compositional setup:
 - Olivine (~50%, refractory)
 - Pyroxene (~35%, fertile)
 - Feldspar (~15%, ^{26}Al)
- Varying solid-melt density contrast, grain sizes (permeability), planetesimal radius, formation time, ...

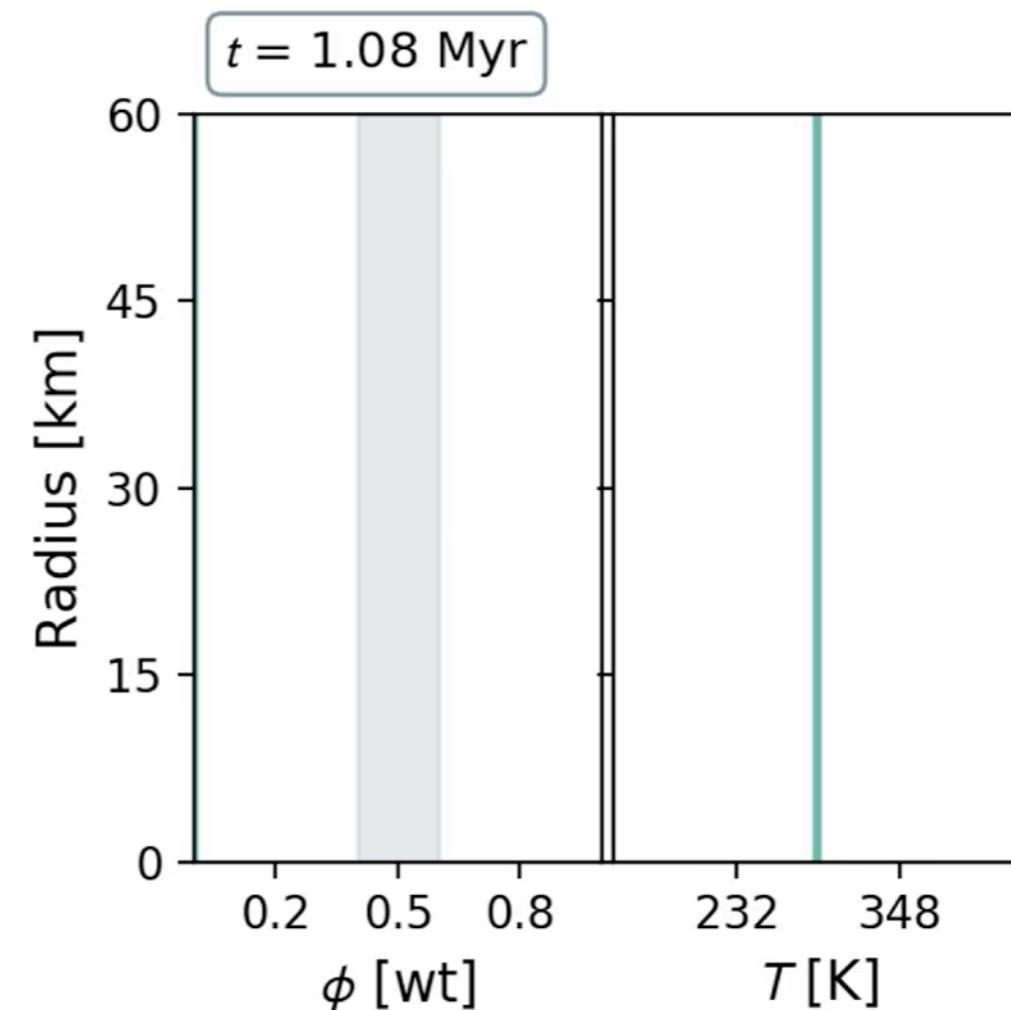
Large-scale melting

‘Magma ocean’



‘Magma sill’

$$R_P = 60 \text{ km}$$



$t_{\text{form}} = 0.72$ Myr, grain size $a = 0.1$ mm

$t_{\text{form}} = 1.08$ Myr, grain size $a = 10$ mm

Melt segregation efficiency

Melt segregation

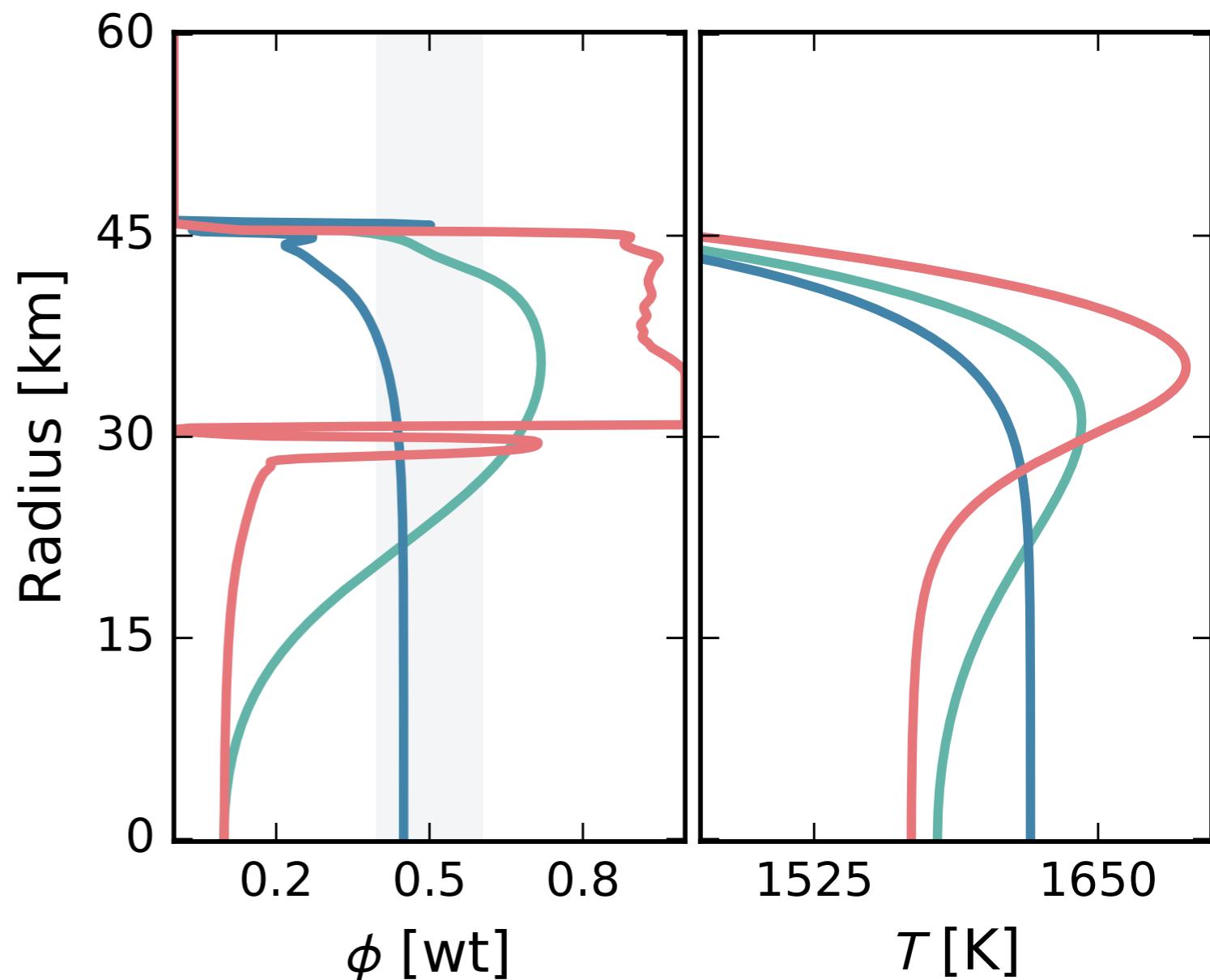
$$\Delta\phi = \phi_{\max} - \phi_{\text{center}}$$

Temperature anomaly

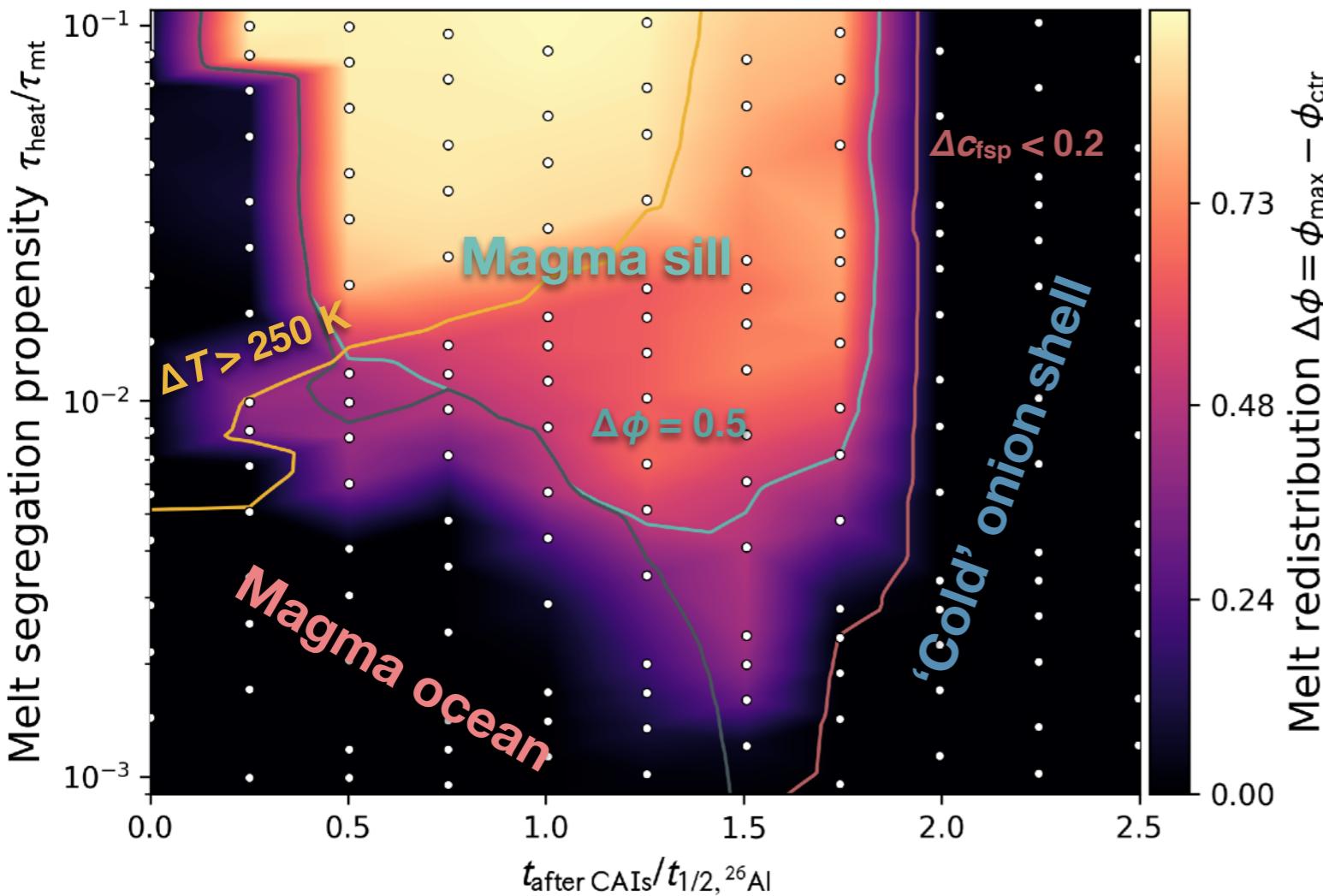
$$\Delta T = T_{\max} - T_{\text{center}}$$

^{26}Al redistribution

$$\Delta C_{\text{fsp}} = C_{\text{fsp},\max} - C_{\text{fsp},\text{center}}$$



Melt segregation regimes



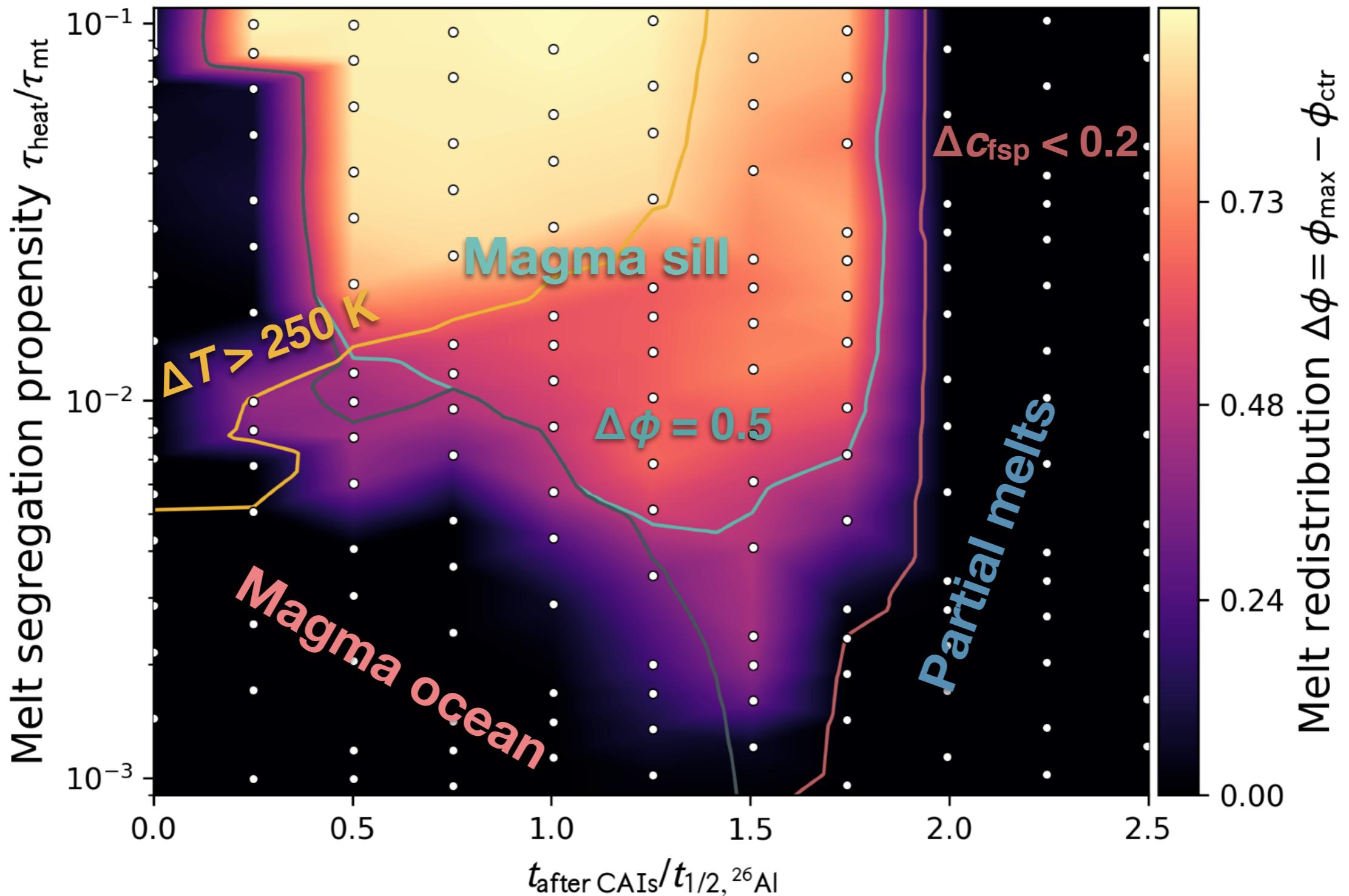
Melting timescale

$$\tau_{\text{heat}} = c_p \cdot \Delta T / H_{\text{Al}}(t)$$

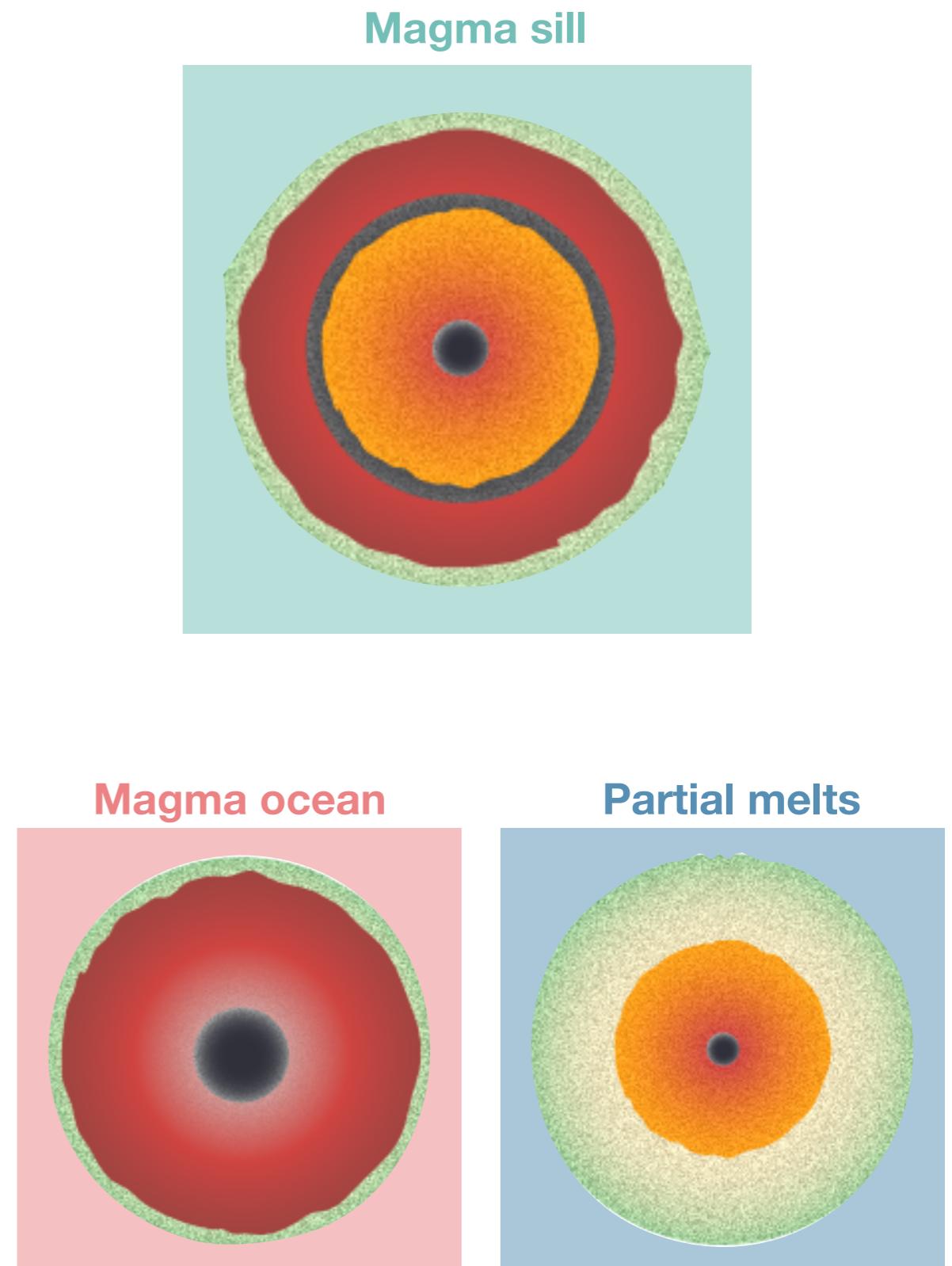
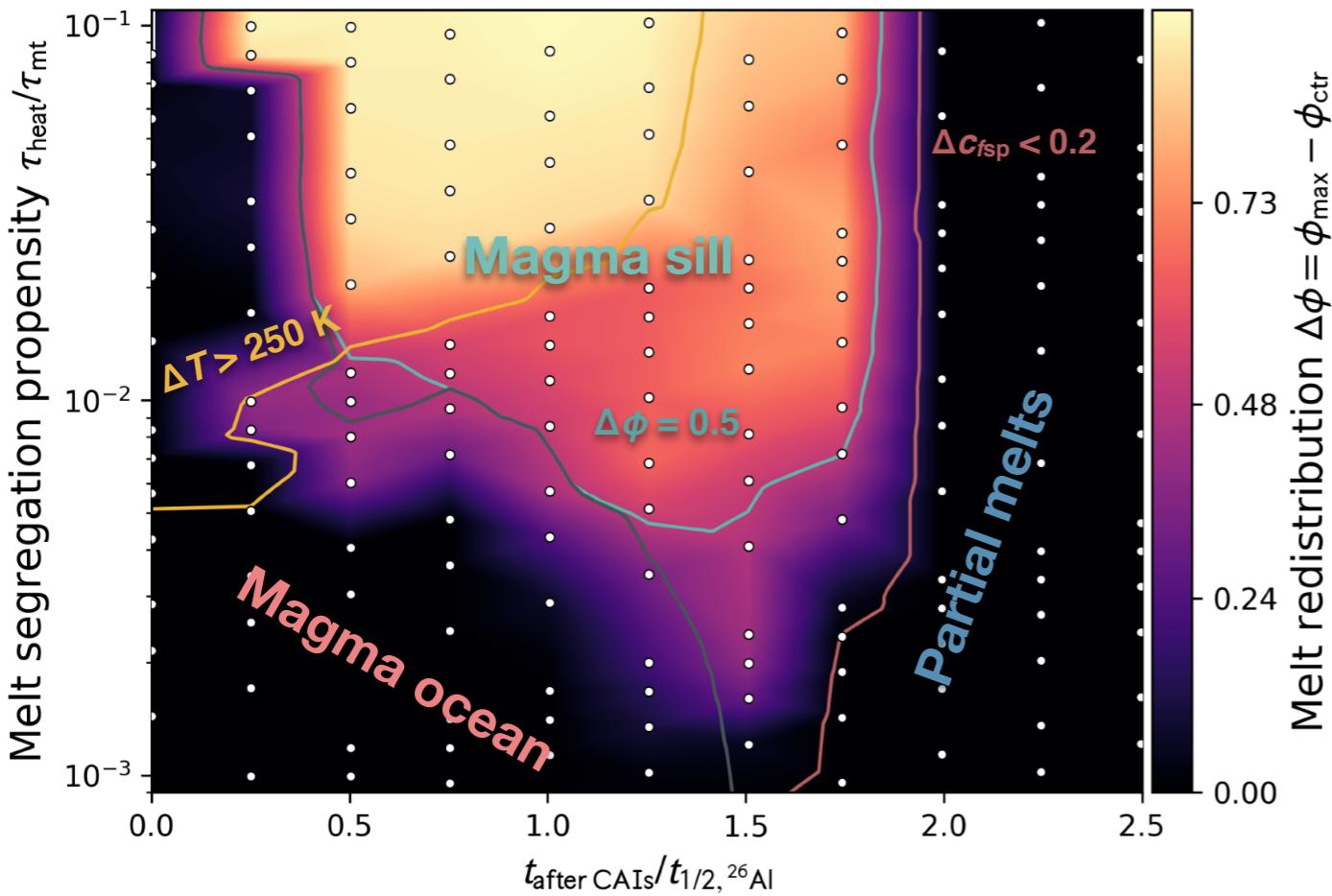
Melt transport timescale

$$\tau_{\text{mt}} = \frac{R_p}{2} / w_s$$

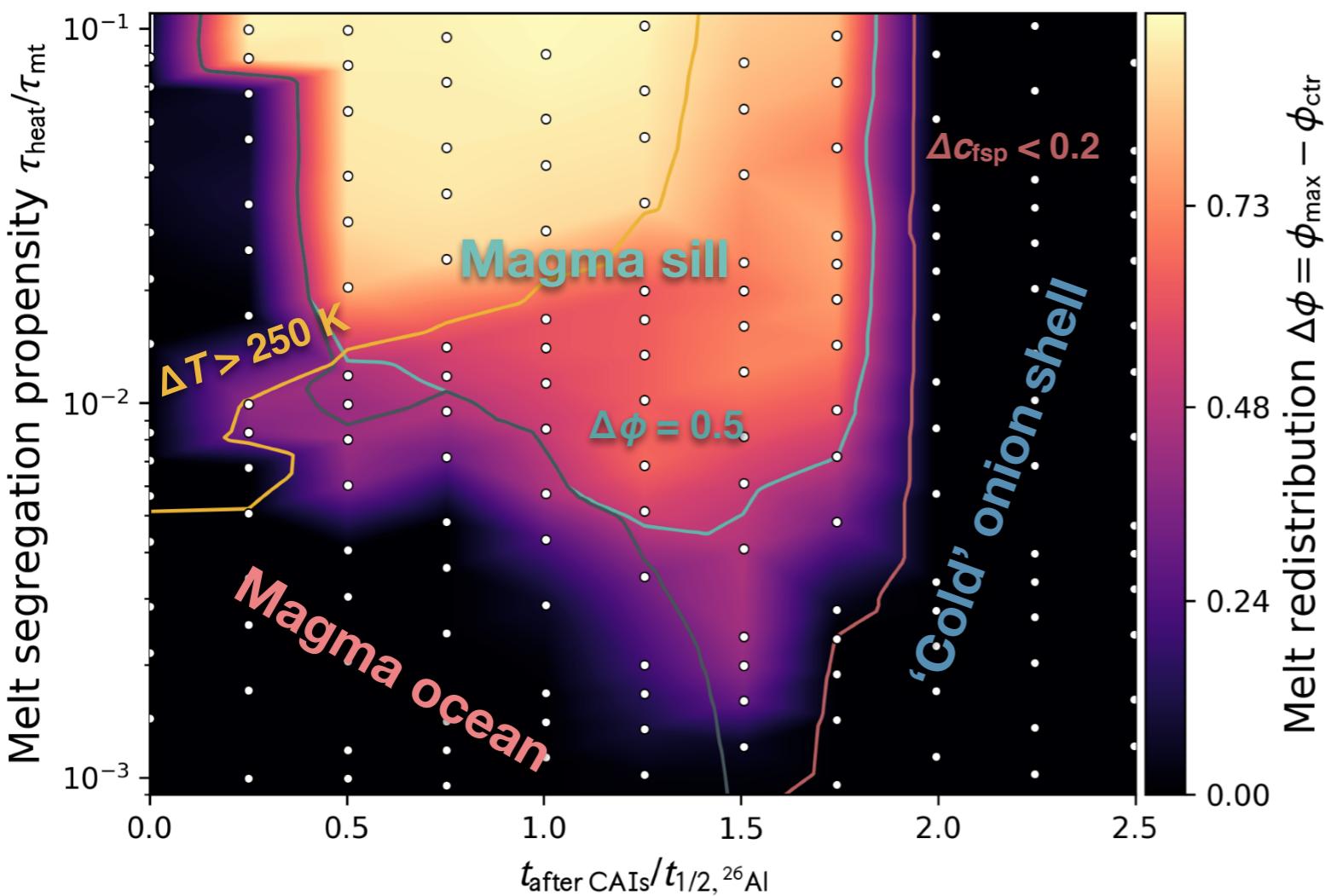
Melt segregation regimes



Summary I – Silicate melt ascension

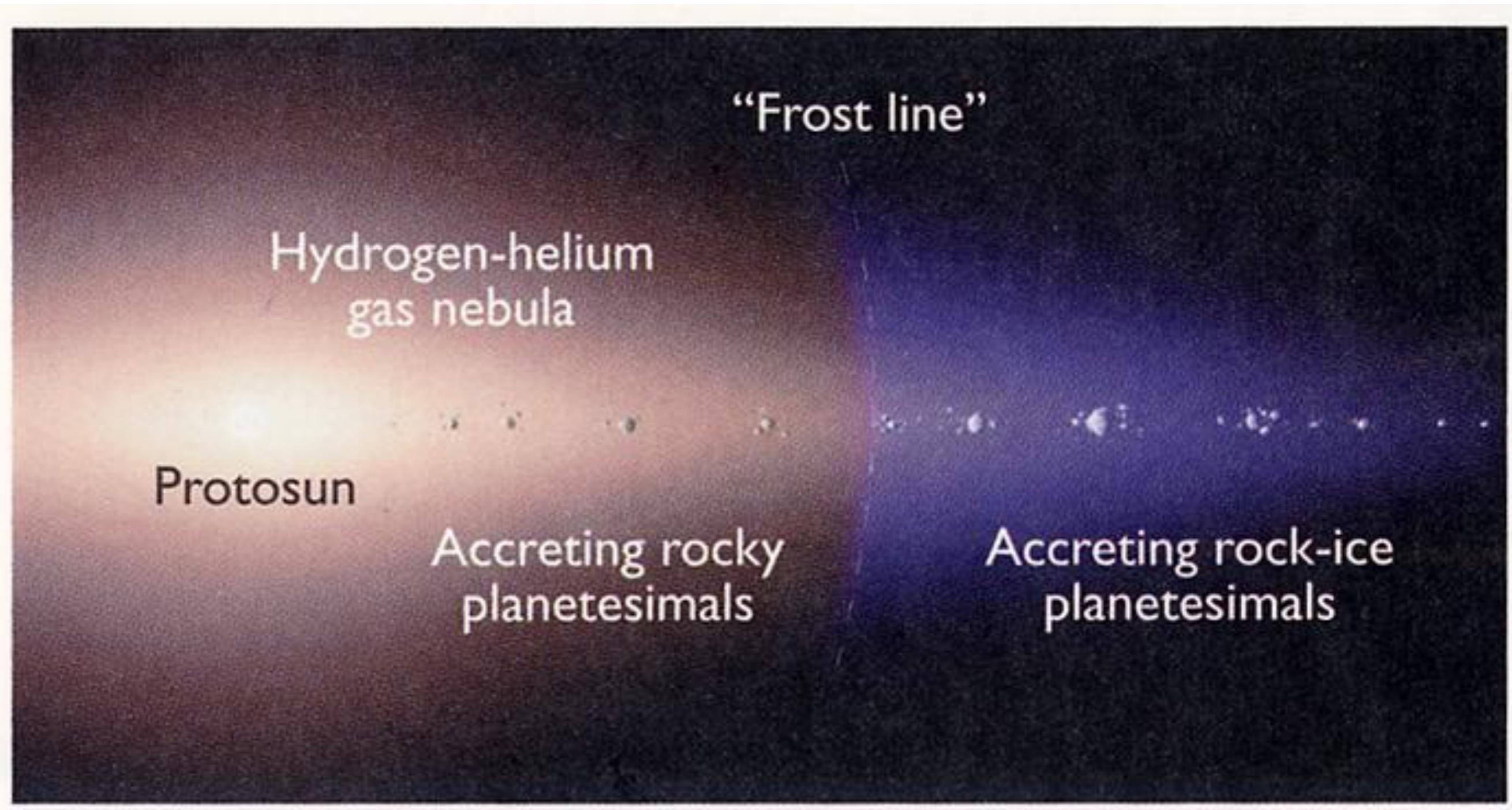


Summary I – Silicate melt ascension

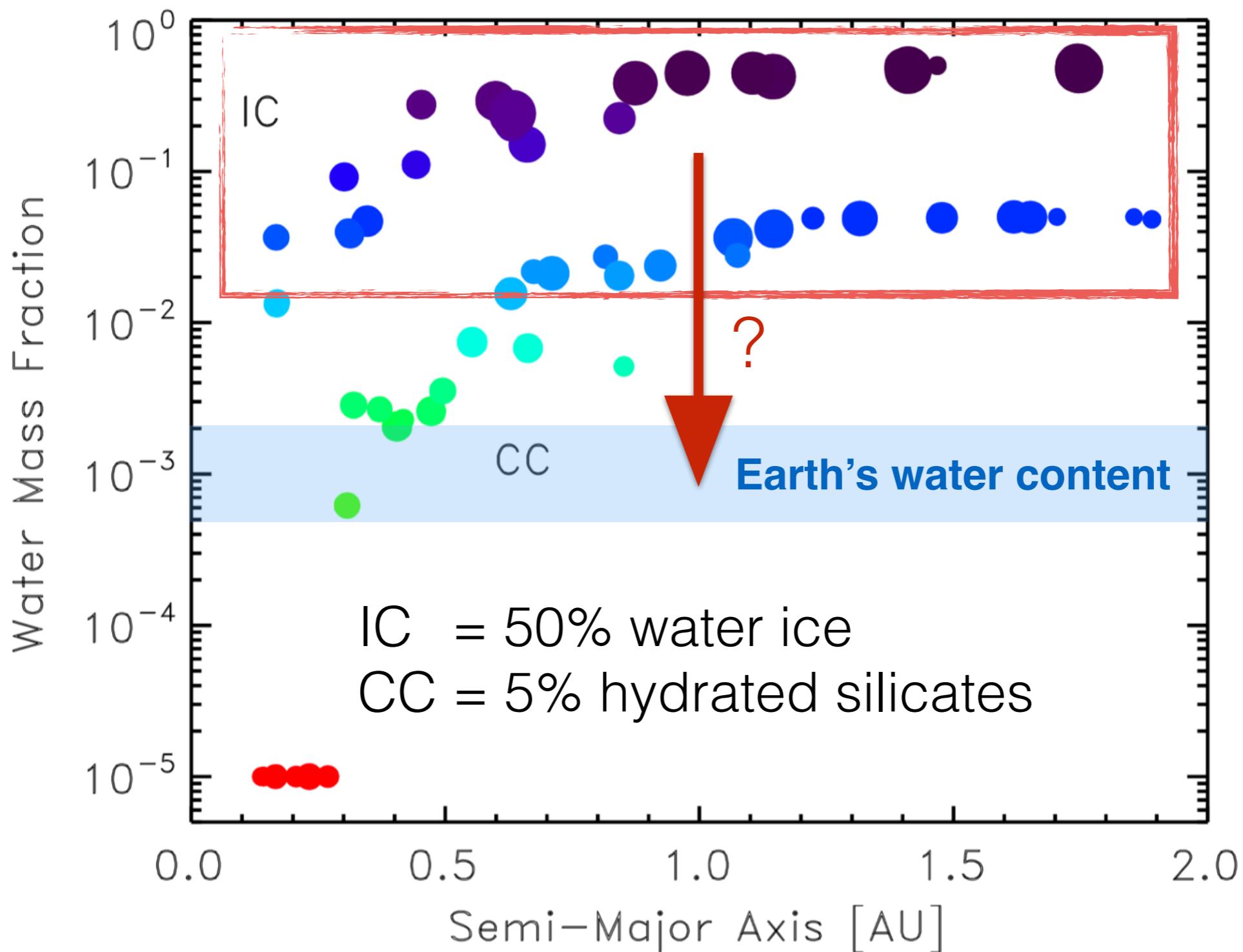


- Melt ascension **crucially dependent on grain size**
- Al partitioning does not generally deplete heat source
 - Structural & compositional state can be related to **evolutionary tracks**

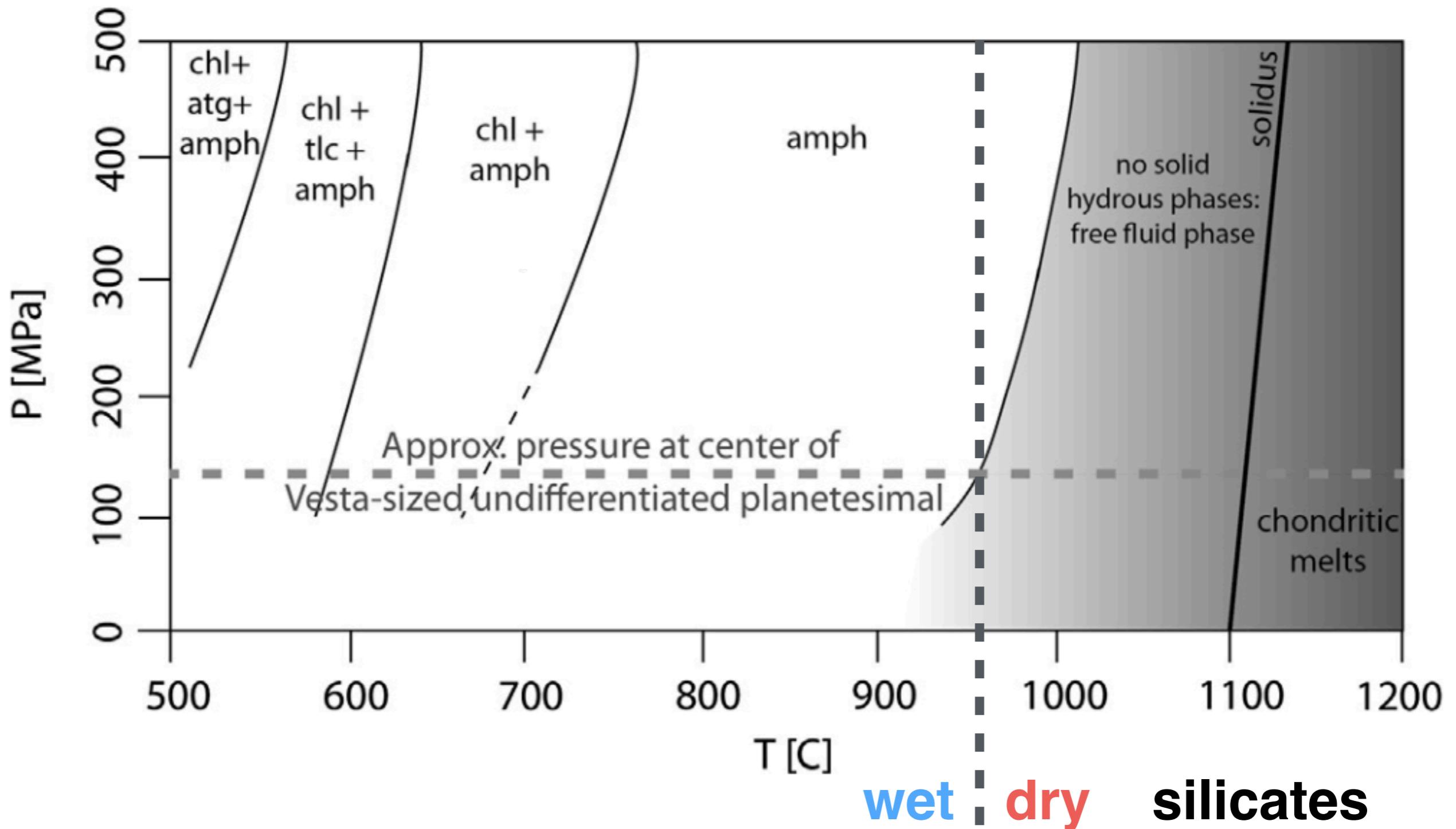
Part II: Dehydration of planetesimals



Fine-tuning Earth-like water abundances

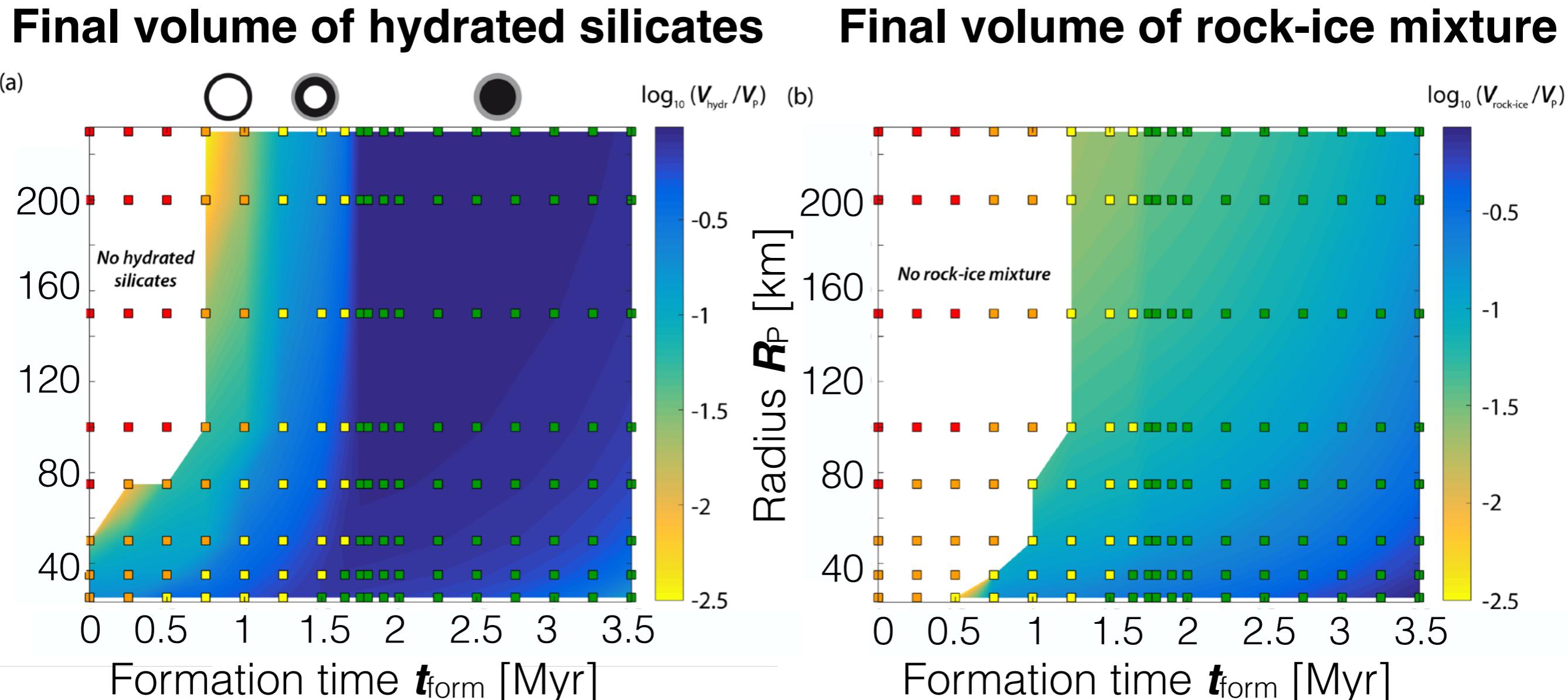


Planetesimal ‘hydrology’



Planetesimal dehydration via outgassing

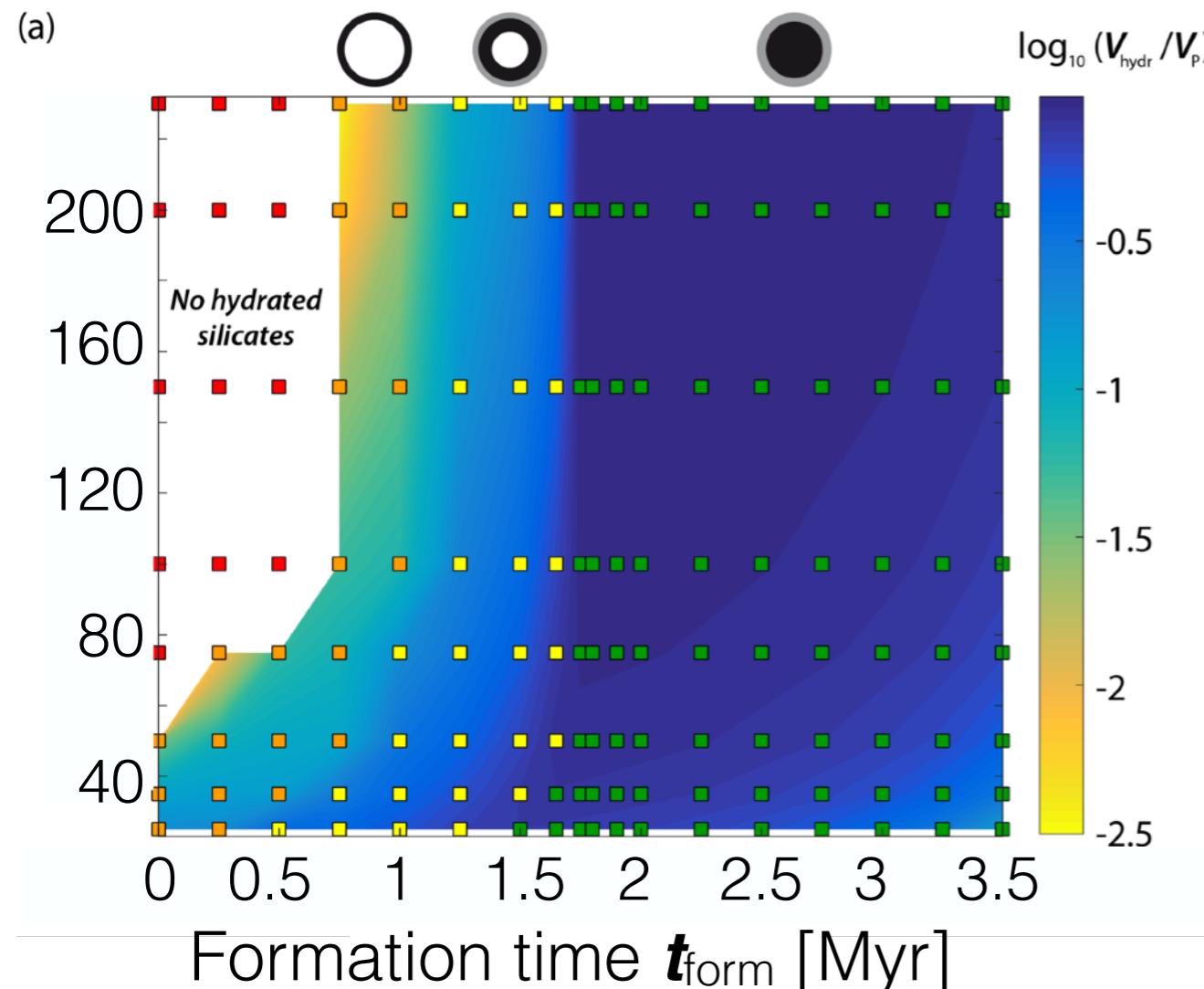
rock-ice mixture hydrated silicates dry



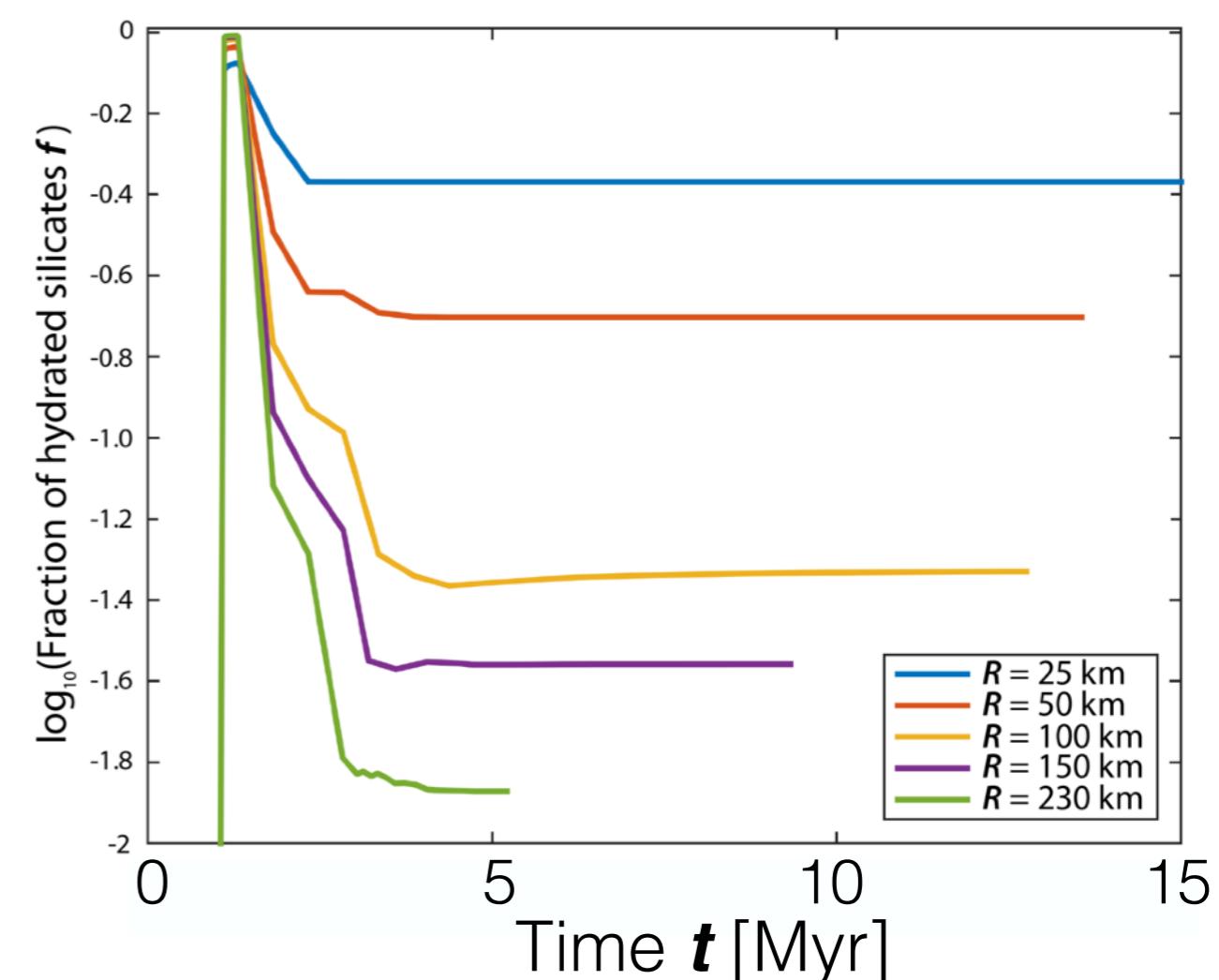
Planetesimal dehydration via outgassing

rock-ice mixture hydrated silicates dry

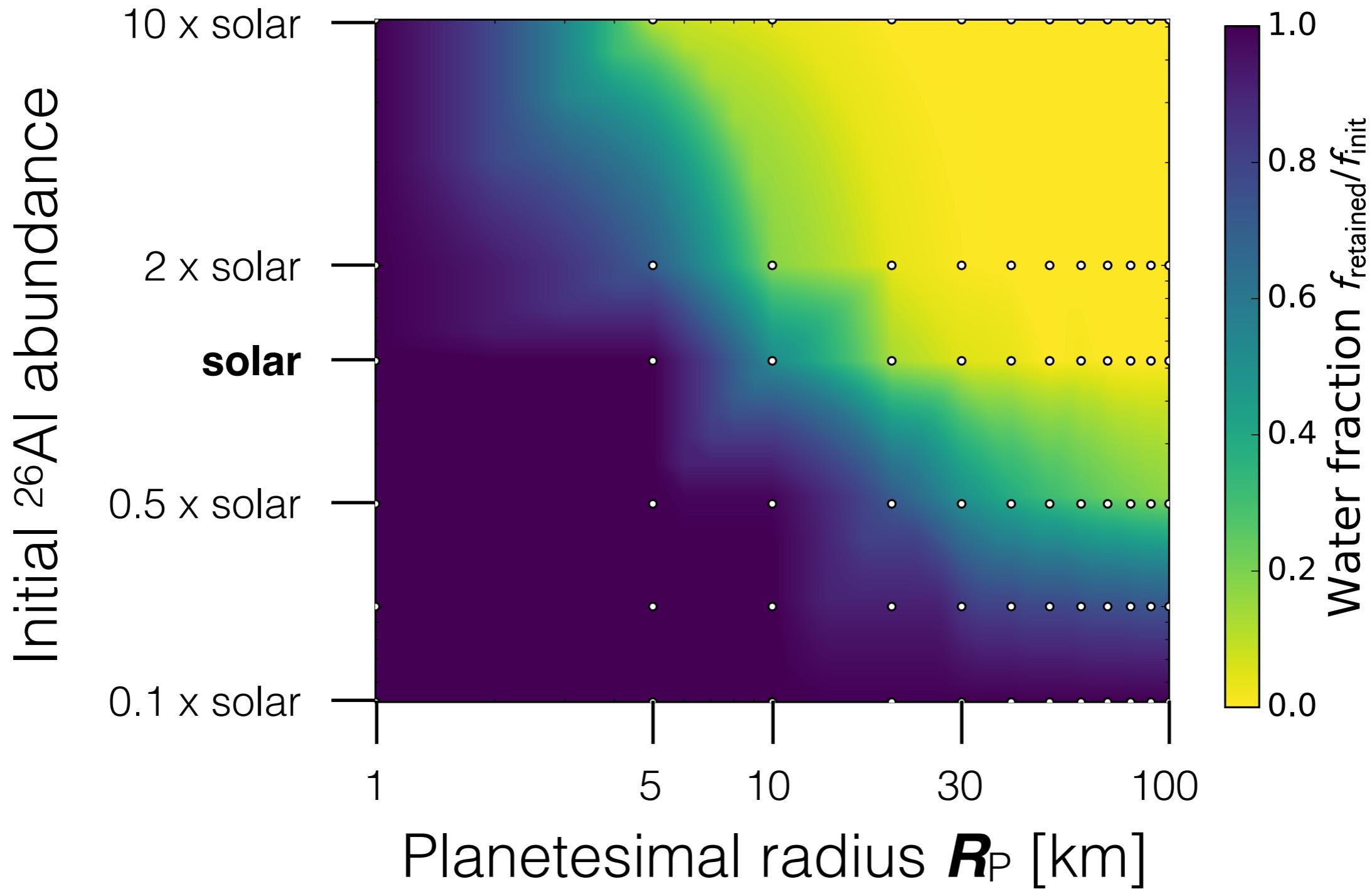
Final volume of hydrated silicates



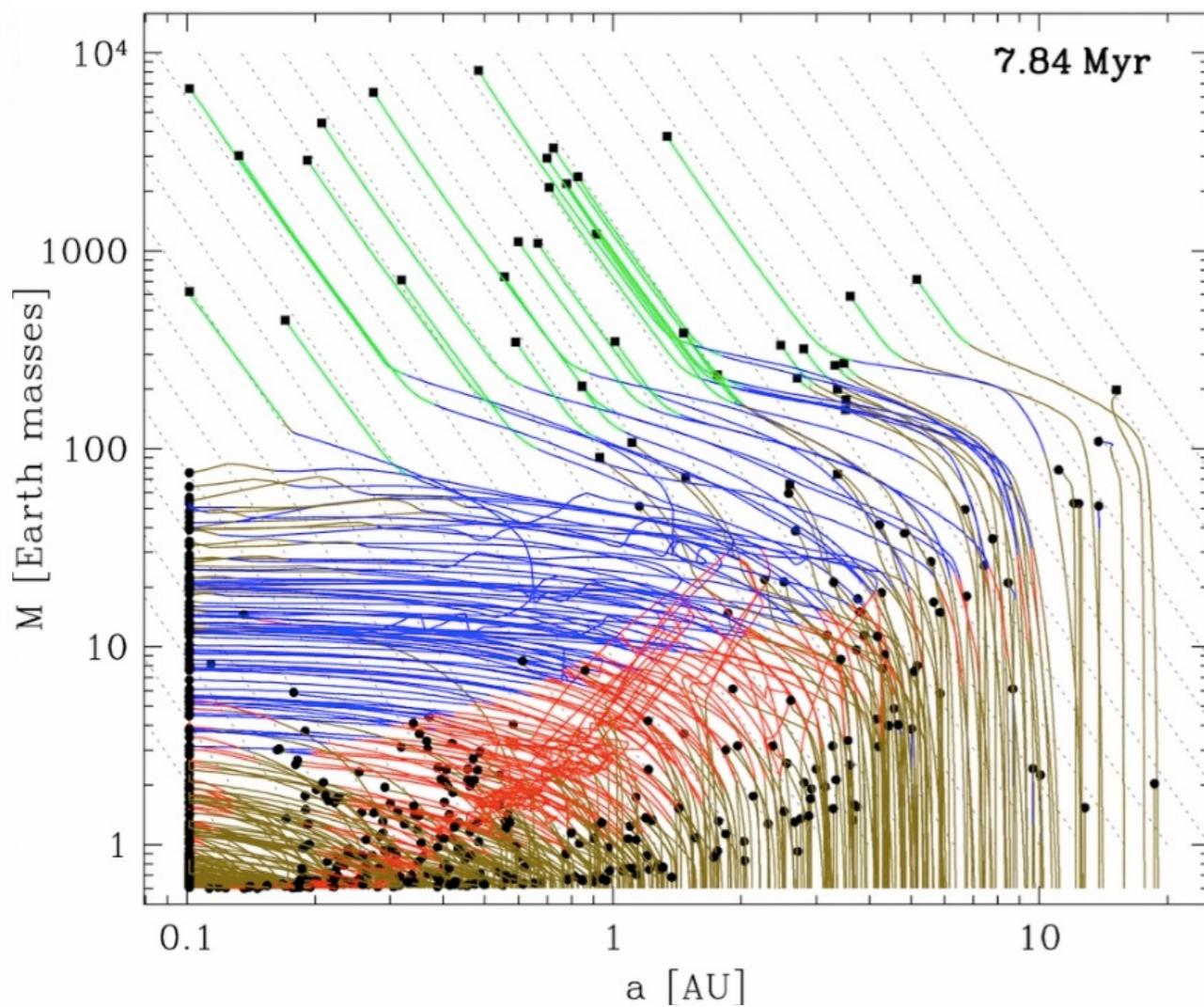
Time evolution



Extrapolation to (exo-)planetary population?

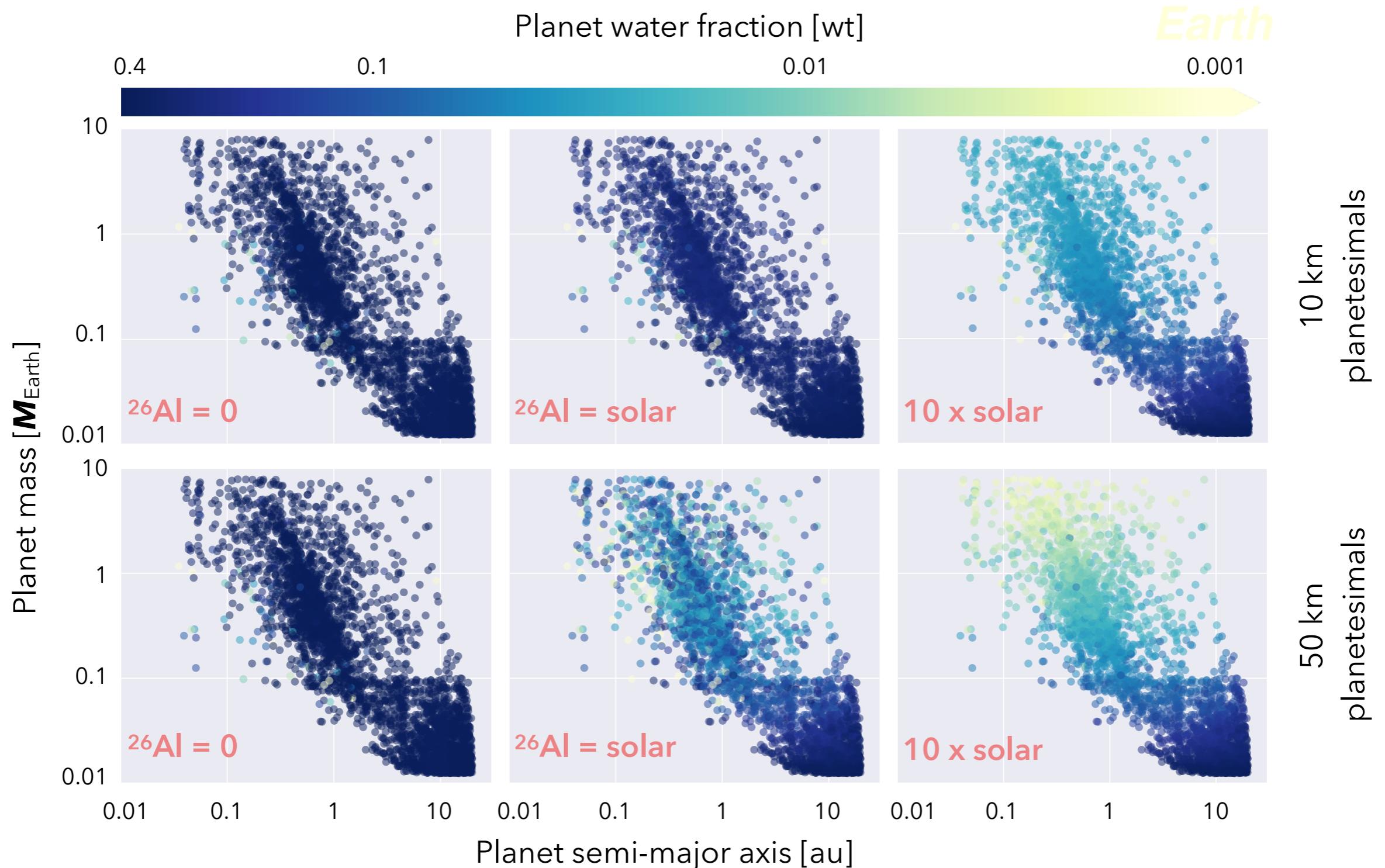


‘Degassing’ planet population synthesis

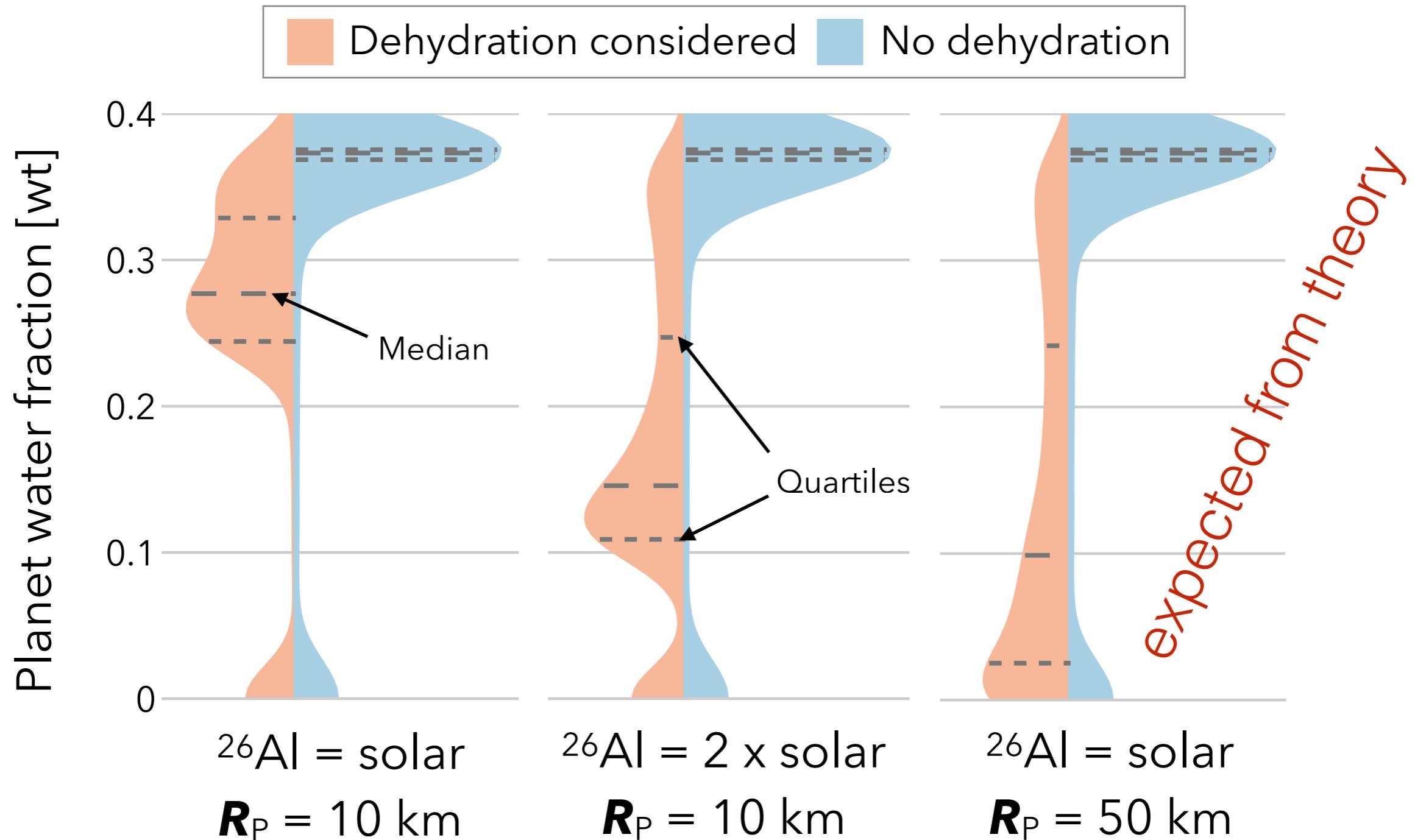


- Planet population synthesis calculations ('Bern model')
- Planetary embryos migrate through ‘sea’ of planetesimals
- Follow planetary built-up and major silicate + volatile phases
- **Add degassing/dehydration from planetesimal models**

'Degassing' planet population synthesis

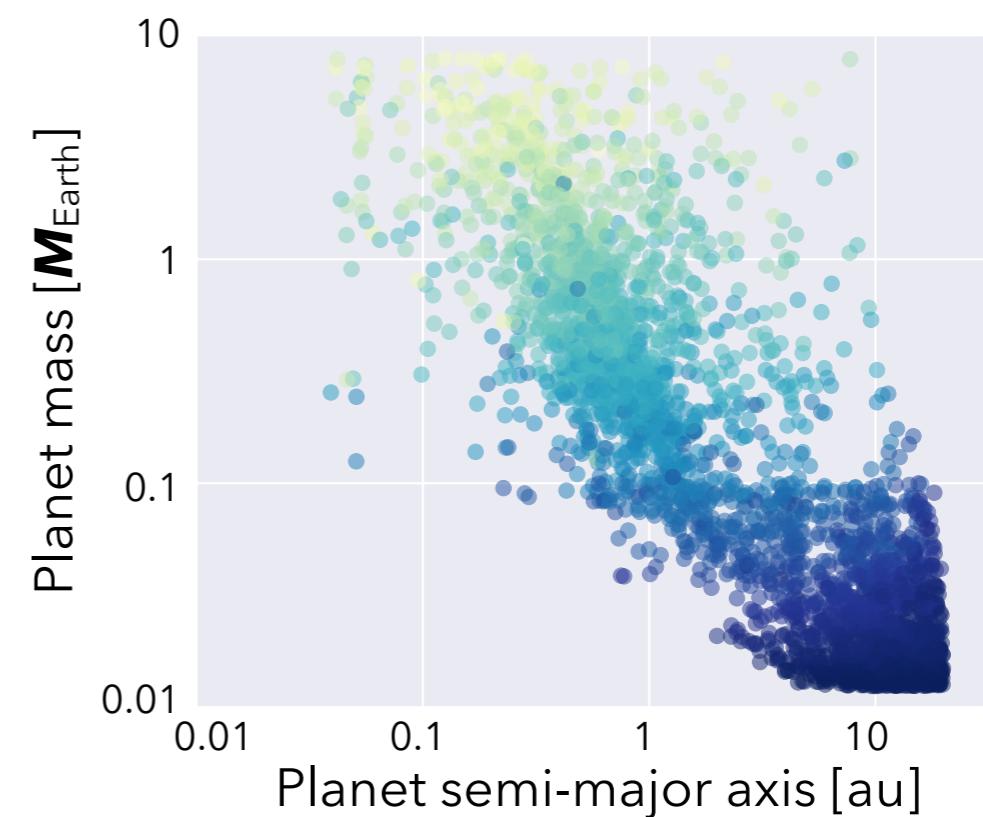


'Degassing' planet population synthesis

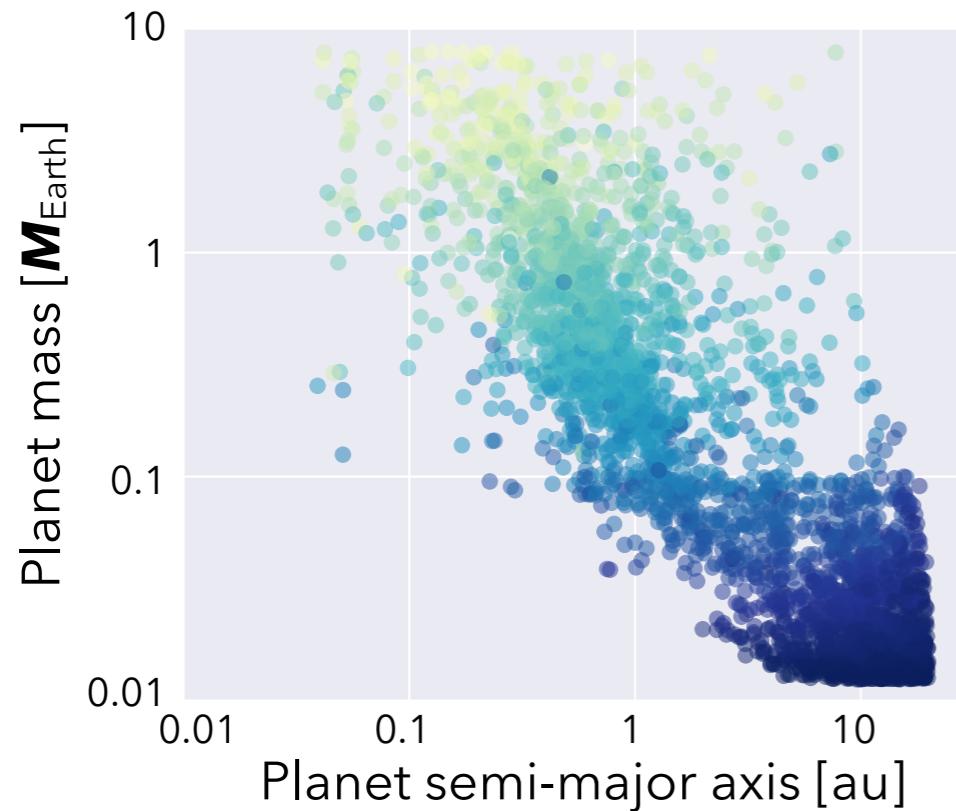
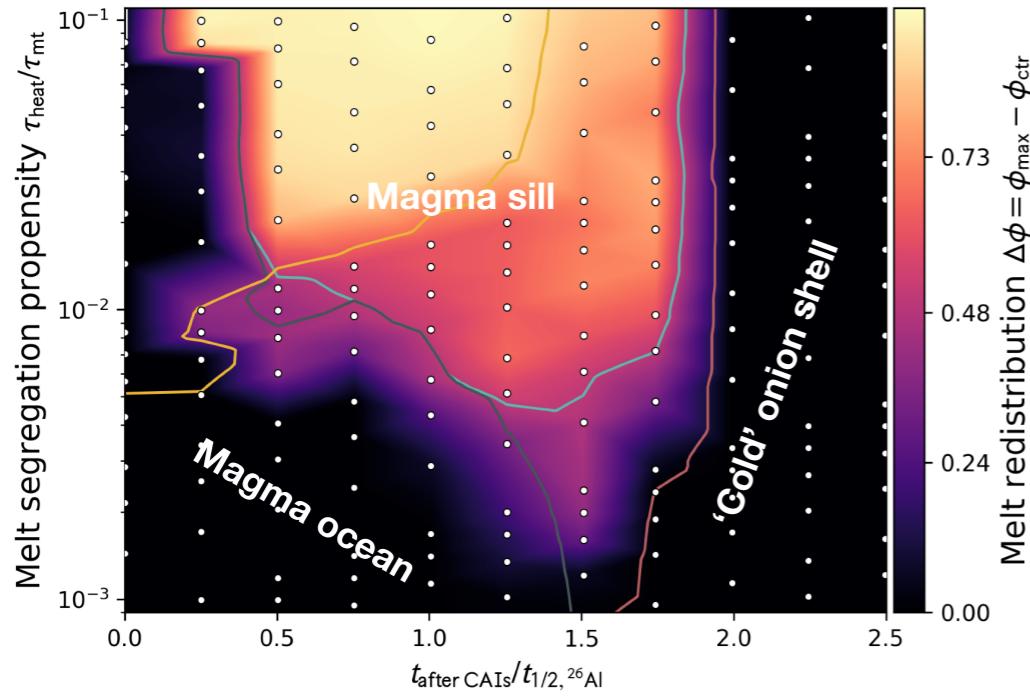


Summary II – Planetesimal devolatilisation

- **Degassing planetesimals** with low ^{26}Al , low-water fraction may be important **contributor of Earth's water**
- **Degassing** processes on planetesimals alter volatile abundances **before accretion, flatten distribution**
- Does ^{26}Al heating imprint **traceable signature** in planetary **volatile abundances?**



Conclusions



1. Interior magma oceans likely existed, **melt segregation only partly effective**
2. Can we identify **primordial planetesimals** left-over from accretion and **assign evolutionary pathways** related to **silicate melt migration**?
3. **Degassing can deplete volatile abundances in terrestrial planets**
 - Signature may be traceable in **exoplanet** census?